



# Interactions between aquatic plants and turbulent flow: A field study using stereoscopic PIV

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A stereoscopic particle image velocimetry (PIV) system for use in shallow ( $\sim 0.5$  m deep) rivers was developed and deployed in the Urie River, Scotland, to study the interactions between turbulent flow and a *Ranunculus penicillatus* plant patch in its native environment. Statistical moments of the velocity field were calculated utilising a new method of reducing the contribution of measurement noise, based on the measurement redundancy inherent to the stereoscopic PIV method. Reynolds normal and shear stresses, their budget terms, and higher order moments of the velocity probability distribution in the wake of the plant patch were found to be dominated by the presence of a free shear layer induced by the plant drag. Plant motion, estimated from the PIV images, was characterised by travelling waves that propagate along the plant with a velocity similar to the eddy convection velocity, suggesting a direct coupling between turbulence and the plant motion. The characteristic frequency of the plant velocity fluctuations ( $\sim 1$  Hz) may suggest that the plant motion is dominated by large eddies with scale similar to the flow depth or plant length. Plant and fluid velocity fluctuations were, in contrast, found to be strongly correlated only over a narrow ( $\sim 30$  mm) elevation range above the top of the plant, supporting a contribution of the shear layer turbulence to the plant motion. Many aspects of flow-aquatic plant interactions remain to be clarified, and the newly developed stereoscopic field PIV system should prove valuable in future studies.

**Key words:** ...

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## 1. Introduction

Aquatic plants play a vital role in the management and healthy functioning of river ecosystems. They provide habitat, refuge, and food for periphyton, invertebrates, and fish; they produce oxygen and sink carbon through photosynthesis; they regulate sediment transport and mixing, and they contribute to hydraulic resistance (e.g. Naden *et al.* 2006; Bornette & Puijalon 2011; Folkard 2011*b*; Nepf 2012). Understanding of these processes is important for the successful management of river systems (mitigating flood risk, preserving biodiversity, maintaining water quality) but is still limited by a lack of fundamental knowledge of the interactions between plants and flowing water. One of the reasons for this is that flow-plant interactions are scale dependant, covering a wide range of scales from the sub-leaf to the plant patch and larger scales, and are thus controlled by several complex and interlinked phenomena such as turbulence, viscous and pressure drag

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39 forces, plant biomechanical properties, and plant motion (Nikora 2010). Another reason  
 40 is that experimental measurement of these phenomena remains challenging. Many of the  
 41 experimental investigations on aspects of flow-plant interactions have been carried out  
 42 in laboratory flumes using artificial plant replicas or plant surrogates (e.g. Ghisalberti &  
 43 Nepf 2002; Nezu & Sanjou 2008; Siniscalchi *et al.* 2012), or using real plants attached to  
 44 the bed in some artificial way (e.g. Sand-Jensen 2003; O’Hare *et al.* 2007; Siniscalchi &  
 45 Nikora 2012). Although these studies allow systematic manipulation of flow conditions  
 46 and deployment of a full array of experimental technologies, it remains an open question  
 47 as to whether they are truly representative of real plants in their natural habitats. A  
 48 number of field studies have also been carried out (e.g. Koehl & Alberte 1988; Sand-  
 49 Jensen & Mebus 1996; Green 2005; Naden *et al.* 2006; Sukhodolova & Sukhodolov 2012).  
 50 These studies, however, inevitably resort to point velocity measurement techniques (of-  
 51 ten involving only time averaged velocities) which miss much of the detailed structure of  
 52 the flow field.

53 In the study reported here, the need for more extensive field data on flow-aquatic  
 54 plant interactions is addressed by developing a stereoscopic PIV system for field use and  
 55 deploying it in the Urie River, Scotland. The PIV technique has been previously used  
 56 outside of a laboratory (e.g. Nimmo Smith *et al.* 2002; Zhu *et al.* 2006; Tritico *et al.*  
 57 2007; Katija & Dabiri 2008; Liao *et al.* 2009). This study, however, is the first time the  
 58 stereoscopic PIV method has been used in the field, allowing all three components of the  
 59 velocity vector to be captured. The system is utilised to study the interactions between  
 60 river turbulence and the motion of a *Ranunculus penicillatus* plant patch in its natural  
 61 environment.

62 The structure of the paper is as follows. First, the design of a stereoscopic PIV system  
 63 that can be deployed in small rivers ( $\sim 0.5$  m flow depth) is discussed, including system  
 64 calibration, analysis algorithms, and a new method of reducing the contribution of mea-  
 65 surement errors to certain velocity statistics. Second, features of the field site selected  
 66 for the study are identified, and measurement errors are analysed. Third, statistics of the  
 67 flow field in the wake of the *Ranunculus* plant patch are evaluated including terms of  
 68 the Reynolds stress budget equation, spectra, and convection velocity. Fourth, statistics  
 69 of the plant motion are evaluated along with correlations between turbulence and plant  
 70 movement. Finally, potential interaction mechanisms between the plants and the flow are  
 71 discussed.

## 72 **2. In-situ stereoscopic PIV system**

73 The in-situ stereoscopic PIV system was designed to utilise existing components from  
 74 a custom-made laboratory PIV system including the laser (Oxford Lasers Nano-L-50/100  
 75 PIV, twin Nd:YAG, 100 mJ at 50 Hz) and cameras (Dalsa 4M60, CMOS,  $2352 \times 1728$   
 76 pixels at 60 frames per second, 7.4 micron pixel pitch, 60% effective fill factor, 532 nm  
 77 bandpass optical filter, 60 mm lens at f/5.6) and direct to disk image recording setup  
 78 ( $4 \times 7200$  rpm SATA disks in RAID 0 per camera). At the core of the design is a glass bot-  
 79 tomed ‘boat’ shaped structure which sits at the water surface and allows a pair of cameras  
 80 and the laser light sheet stable optical access through the fluctuating water surface of the  
 81 river (figure 1). The streamlined design of the ‘boat’ limits the disturbed region of the  
 82 flow-field to a thin boundary layer near the water surface estimated to be approximately  
 83 5 mm thick (based on previous experience with similar structures and approximate es-  
 84 timates using conventional relationships). The ‘boat’ incorporates a trapezoidal shaped  
 85 water prism (e.g. Prasad 2000) to minimise both optical distortion caused by refraction  
 86 and internal reflections that occur at the water-glass-air interfaces. The ‘boat’, cameras,

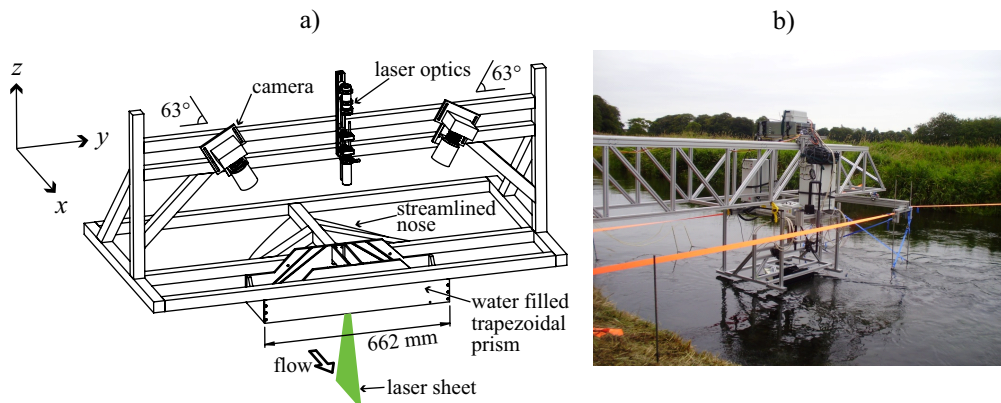


FIGURE 1. Schematic of glass bottomed ‘boat’ attached to camera and laser mount sub-frame (a). System deployed in the Urie River, Scotland (b).

87 and laser optics sit on a rigid sub-frame that allows the cameras and laser optics to be  
 88 aligned, focussed and calibrated in the laboratory prior to field deployment. In the field,  
 89 the sub-frame is attached to a specially designed frame (bridge) and carriage assembly  
 90 that allows the PIV system to be traversed in the streamwise (0.5 m) and transverse  
 91 (5.0 m) directions. The bridge is constructed of aluminium extrusions (Kanya PVS), it  
 92 spans 7.5 m, weighs 150 kg, and is designed so that at least one end of the bridge is  
 93 anchored on the river bank. The other end of the bridge can be supported mid river  
 94 on stainless steel poles with tension straps tied to the far river bank to ensure stability.  
 95 The turbulent wakes created by the bridge support elements are well clear of the mea-  
 96 surement area (figure 1b). The laser, cameras and computer are powered by a portable  
 97 5 kVA generator. Seeding (conifer pollen, 60-80 micron diameter, 800-1000 kg/m<sup>3</sup> den-  
 98 sity) is mixed with water at a concentration of 100 g/l and injected into the river by a  
 99 pump approximately 5 m upstream of the test section at a solids rate of 100 grams per  
 100 minute. The entire bridge, carriage, and laser and camera assembly can be installed at a  
 101 field site by an 8 person team in around 7 hours. Disassembly is faster (around 2 hours)  
 102 leaving several hours for measurements during a single day deployment. The orientation  
 103 of our coordinate system is shown in figure 1. We will refer to the  $x$ ,  $y$ , and  $z$  axis and  
 104 their associated velocity components  $u$ ,  $v$ , and  $w$  as the nominal streamwise, transverse  
 105 and bed normal (or vertical) directions and velocities, respectively. In practice, the laser  
 106 light sheet was aligned visually to be parallel to the local mean flow direction by making  
 107 use of the visible stream of tracer particles injected upstream.

108 A stereoscopic camera configuration was selected because it offers a number of bene-  
 109 fits over a single orthogonal camera setup. Firstly, all three components of the velocity  
 110 vector are resolved compared to just two components for a single camera configuration.  
 111 The additional velocity component provides valuable information on the structure and  
 112 dynamics of the flow field, particularly in the highly three-dimensional flow regions near  
 113 the bed of open channels and in the wake of aquatic plants. Secondly, the stereoscopic  
 114 configuration allows all cameras and laser optics to be placed above the river surface.  
 115 This minimises the disturbance to the flow field and limits camera vibration which can  
 116 introduce additional error into the velocity measurements. The need to waterproof the  
 117 camera and laser components is also removed. Thirdly, perspective errors which occur  
 118 in single camera systems due to the unresolved out of plane velocity component (Raffel  
 119 *et al.* 2007) are eliminated by the stereoscopic configuration. Finally, by taking advantage  
 120 of the redundancy inherent in stereoscopic PIV, some velocity statistics can be calculated

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121 with a significantly reduced contribution of random measurement noise. In the following  
 122 section, details of our implementation of the stereoscopic PIV method are outlined, in-  
 123 cluding: calibration and stereoscopic reconstruction, cross correlation algorithms, method  
 124 of extracting the velocity of plant motion, and analysis of measurement errors.

### 125 2.1. Stereoscopic PIV calibration

126 Our stereoscopic PIV implementation is based on the ‘mapping’ method introduced by  
 127 Willert (1997), where cross correlation is performed on images that have been ‘dewarped’  
 128 to obtain a constant magnification across the image. The 2-component vector fields from  
 129 a pair of cameras are subsequently combined to reconstruct the three-component velocity  
 130 field. Critical to both the image dewarping and velocity field reconstruction steps is a  
 131 function which relates three-dimensional ( $x, y, z$  - streamwise, transverse, bed-normal re-  
 132 spectively) ‘world’ coordinates to corresponding two-dimensional image coordinates. To  
 133 obtain it, we use a pinhole camera model (e.g. Calluaud & David 2004) combined with a  
 134 2-media refraction model (neglecting the contribution of the glass elements of the water  
 135 prism) based on Maas (1996) and a misalignment correction based on Wieneke (2005).  
 136 In total, 13 model parameters need to be estimated for each camera using a calibra-  
 137 tion procedure, including four intrinsic camera parameters ( $f_x, f_y, i_0, j_0$ ), six extrinsic  
 138 camera parameters ( $\alpha, \beta, \gamma, t_x, t_y, t_z$ ), and three parameters for the refraction model  
 139 ( $\alpha_g, \beta_g, t_{zg}$ ). Three additional parameters ( $\alpha_m, \beta_m, t_m$ ) apply to all cameras and are  
 140 used to correct any misalignment between the laser light sheet and the calibration target.  
 141 Here  $f_x$  and  $f_y$  are camera focal lengths,  $i_0$  and  $j_0$  are image origin coordinates,  $\alpha, \beta, \gamma$   
 142 and  $t_x, t_y, t_z$  are the three Euler rotation angles and three translations, respectively, that  
 143 define the position and viewing direction of the camera. The refraction model param-  
 144 eters  $\alpha_g, \beta_g$  and  $t_{zg}$  are two rotation angles and one translation that give the position  
 145 and orientation of the water-air interface, while misalignment parameters  $\alpha_m, \beta_m$  and  
 146  $t_m$  map the light sheet plane onto the calibration plane. Other parameters are available  
 147 to incorporate lens distortions or to incorporate the refraction caused by the glass win-  
 148 dows of the ‘prism boat’, but these parameters are not used in this study as they were  
 149 found to not improve the calibration. The calibration procedure is carried out in a lab-  
 150 oratory tank after final alignment and focussing of the cameras. The intrinsic, extrinsic,  
 151 and refraction parameters are estimated for each camera based on a set of images of a  
 152 two-sided calibration plate (3 mm diameter dots spaced at 20 mm) which is translated  
 153 to different positions using a precision machined baseplate. The calibration images pro-  
 154 vide a set of point coordinates (the centres of each dot on the calibration plate image)  
 155 and corresponding world coordinates (based on the known calibration plate geometry)  
 156 allowing the model parameters to be optimised using an iterative least square fit. Finally,  
 157 the misalignment correction parameters are estimated from the experimental PIV images  
 158 by ensemble cross correlation between the dewarped images from the first and second  
 159 cameras (Wieneke 2005). In this way, the precise position of the light sheet relative to  
 160 the cameras does not need to be fixed in the laboratory and some adjustment in the field  
 161 is possible (as long as the light sheet remains within the camera depth of field).

### 162 2.2. Cross correlation algorithm

163 A detailed description of our cross correlation algorithm and evaluation of its performance  
 164 is available in Cameron (2011). Some modifications were necessary to optimise for the  
 165 field PIV images; these are described in this section.

166 Our PIV algorithm can be classed as an iterative deformation method (IDM) with  
 167 windowed Fourier transform based cross correlation. Two key features of the algorithm,  
 168 which directly influence measurement noise, measurement resolution, number of outliers,

169 and the number of iterations required to reach convergence are 1) the size and weighting  
 170 of the interrogation regions (image subsections used for cross correlation analysis, e.g.  
 171 Raffel *et al.* 2007), and 2) the low pass filtering of the velocity field after each iteration. To  
 172 analyse the field PIV images, which have a scale factor of 12 pixels/mm, we have selected  
 173 Blackman weighted 96×96 pixel (8×8 mm) interrogation regions (BL96) with a 12 pixel  
 174 (1 mm) grid spacing, and a low pass filter based on a windowed sinc function (Sinc2.5).  
 175 The modulation transfer function (*MTF*) for this algorithm (IDM-BL96-Sinc2.5) has  
 176 been estimated following Astarita (2007) and is given in figure 2. The *MTF* reflects  
 177 the spatial averaging (low-pass filtering) of the velocity field associated with the cross  
 178 correlation algorithm. For a *MTF* value of 0.9, figure 2 indicates the cut-off wavelength  
 179 (resolution) for IDM-BL96-Sinc2.5 along the  $k_x = 1/\lambda_x$  wavenumber axis is 92 pixels  
 180 (7.7 mm), where  $k_x$  is the wavenumber and  $\lambda_x$  is the wavelength in the streamwise di-  
 181 rection. This algorithm trades in some resolution relative to IDM-BL64-TH6 (Cameron  
 182 2011, figure 2) in return for improved robustness against outliers due to the larger in-  
 183 terrogation regions. In comparison to the classic PIV method with 32 pixel unweighted  
 184 interrogation regions (IDS-TH32, figure 2), IDM-BL96-Sinc2.5 has slightly increased res-  
 185 olution, improved flatness in the pass band, efficient anti-aliasing due to steep roll off  
 186 and negligible side lobes, and significantly increased robustness due to having nine times  
 187 more pixels in each interrogation region. Theoretical convergence for IDM-BL96-Sinc2.5  
 188 is eight iterations, defined here as the number of iterations required for the equivalent  
 189 noise bandwidth ( $ENBW_q$ ) to reach 99.9% of its ultimate value.  $ENBW_q$  is calculated  
 190 for each iteration ( $q$ ) by integrating the squared transfer function predicted after each  
 191 iteration:

$$ENBW_q = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} [MTF_q(k_x, k_z)]^2 dk_x dk_z \quad (2.1)$$

192 where  $MTF_q$  is the modulation transfer function estimated for  $q$  iterations of the algo-  
 193 rithm using the method of Astarita (2007), and  $k_z$  is the wavenumber in the vertical  
 194 direction. For comparison, theoretical convergence of IDM-BL64-TH6 is 36 iterations.  
 195 There is additional filtering of the velocity field associated with the finite thickness of  
 196 the light sheet ( $\sim 1.5$  mm), but in this case the light sheet is quite thin and the effect is  
 197 small relative to the filtering associated with the cross correlation algorithm. The effect  
 198 of the finite resolution of the measurement system is to reduce the contribution of high  
 199 wavenumber (small) eddies to the measured velocity variance. The magnitude of this  
 200 effect depends on the flow field and cannot be easily quantified. For the present experi-  
 201 ments, however, the cut-off wavelength of the measurements (7.7 mm) is small compared  
 202 to the flow depth (390 mm) and therefore it is likely that the missing velocity variance  
 203 is small.

204 A feature of the field PIV images is that some of the interrogation regions were inter-  
 205 mittently occupied by plant material or void of sufficient seeding particles such that valid  
 206 velocity vectors could not be obtained. In order to pre-empt these problems we introduce  
 207 a measurement ‘clipping’ function,  $\phi_M(x, z, t_n)$ , ( $t_n$  is the time step) defined as  $\phi_M = 1$   
 208 for valid interrogation regions (with sufficient seeding and absent of any plant), otherwise  
 209  $\phi_M = 0$ . The mean value of  $\phi_M(x, z, t_n)$  can be defined as the measurement porosity.  
 210 The measurement clipping function serves two purposes, first it is passed to the cross  
 211 correlation algorithm so that bad interrogation regions can be handled appropriately by  
 212 the algorithm, and second it is passed to velocity field post processing routines so that  
 213 velocity statistics are correctly calculated only over valid data. Regions where  $\phi_M = 0$   
 214 are identified in an image pre-processing stage using a type of signal to noise ratio. Each  
 215 image is first decomposed into two parts, a plant image (by applying a median filter to

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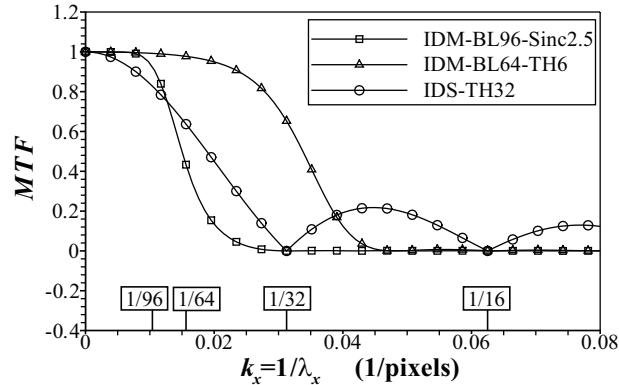


FIGURE 2. Comparison of transfer functions for different PIV algorithms: IDM-BL96-Sinc2.5 is employed in this study; other algorithms are discussed in Cameron (2011).

216 the original image) and a seeding image (by subtracting the plant image from the original  
 217 image). The ‘signal’ is then defined as the sum of the pixel intensities within an interroga-  
 218 tion region for the seeding image, and the ‘noise’ is obtained as the sum of pixel intensities  
 219 within an interrogation region for the plant image. For each interrogation region, if the  
 220 signal divided by the noise is above a threshold value,  $\phi_M$  is set to 1, otherwise it is 0.  
 221 The threshold value was optimised by visual assessment and a trial and error approach  
 222 on a subset of the PIV images. Once determined, it was applied globally throughout the  
 223 image set. Within the PIV algorithm, regions of  $\phi_M = 0$  are replaced with interpolated  
 224 or extrapolated velocity values from neighbouring valid data. This step is important with  
 225 IDM PIV algorithms as it allows vector field low-pass filtering and interpolation to be  
 226 performed in each iteration without inadvertently propagating bad vectors to adjacent  
 227 interrogation regions. We emphasise that the interpolated/extrapolated values are only  
 228 used within the cross correlation algorithm. Time averaged statistics of the velocity field  
 229 are calculated incorporating only valid data as, for example, in the case of the first order  
 230 statistics:

$$\bar{\theta}(x, z) = \frac{1}{\sum_{t_n=1}^{t_n=T} [\phi_M(x, z, t_n)]} \sum_{t_n=1}^{t_n=T} [\theta(x, z, t_n) \times \phi_M(x, z, t_n)] \quad (2.2)$$

231 where  $\theta$  is a flow variable,  $t_n$  is the time step, and  $T$  is the total number of time steps. The  
 232  $\phi_M$  parameter is also used in calculating correlation functions and spectra; equations are  
 233 given where appropriate in the following sections. Potential measurement uncertainties  
 234 associated with flow regions having small values of measurement porosity are limited by  
 235 only presenting data for which  $\bar{\phi}_M > 0.75$ .

236

### 2.3. Calculating plant velocity

237 The field PIV images contained enough detail of the fluctuating plant to extract estimates  
 238 of the vertical  $w_p(x)$  and transverse  $v_p(x)$  plant velocity components. We selected rect-  
 239 angular interrogation regions ( $96 \times 1024$  pixels,  $8 \times 85$  mm) which were sufficiently high  
 240 to cover the entire visible plant cross section, but narrow enough that plant velocities  
 241 could be measured as a function of streamwise position. Standard cross correlation ap-  
 242 plied to median filtered PIV images (to remove seeding particles) resulted in a rather  
 243 wide peak in the correlation function (proportional to the width of the plant) and there-  
 244 fore poor accuracy in estimating the displacement. To improve cross correlation perfor-  
 245 mance, we have employed Wernet’s (2005) symmetric phase only filtering (SPOF) which

246 makes the correlation function more sensitive to the high wave number content of the  
 247 image (i.e., the sharply defined edges of the plant stems and leaves). The SPOF takes  
 248 the form of a weighting function  $C(k_m, k_n)$  (2.3) applied to the cross spectral density  
 249  $G(k_m, k_n) H^*(k_m, k_n)$  between a pair of interrogation regions, where  $G$  and  $H$  are the  
 250 Fourier transforms of the first and second interrogation regions,  $H^*$  is the complex con-  
 251 jugate of  $H$ , and  $k_m$  and  $k_n$  are wavenumbers in the  $m$  and  $n$  dewarped image directions.  
 252 The cross correlation function  $\varphi$  (2.4) is then calculated as the inverse Fourier transform  
 253 ( $FFT^{-1}$ ) of the weighted cross spectral density, which for  $C = 1$  is the standard Fourier  
 254 transform based cross correlation:

$$C(k_m, k_n) = \frac{1}{\sqrt{|G(k_m, k_n)|} \sqrt{|H(k_m, k_n)|}} \quad (2.3)$$

255

$$\varphi = FFT^{-1} [C(k_m, k_n) G(k_m, k_n) H^*(k_m, k_n)] \quad (2.4)$$

256 The displacement of the correlation peak is estimated only in the vertical ( $n$ ) image  
 257 direction which is sensitive to vertical and transverse displacements of the plant. By  
 258 combining the displacements estimated from a pair of stereoscopic cameras, the vertical  
 259  $w_p(x)$  and transverse  $v_p(x)$  plant velocity components are recovered. Due to the extended  
 260 interrogation regions, the measured velocities approximate the cross-sectional average of  
 261 plant velocity fluctuations.

262

#### 2.4. Stereoscopic velocity field reconstruction

263 Two-component velocity fields estimated using cross correlation on dewarped images are  
 264 combined from two cameras to reconstruct the three-component velocity field according  
 265 to Raffel *et al.* (2007) as:

$$\begin{bmatrix} 1 & 0 & \psi_{1m} \\ 0 & 1 & \psi_{1n} \\ 1 & 0 & \psi_{2m} \\ 0 & 1 & \psi_{2n} \end{bmatrix} \begin{bmatrix} \Delta_x \\ \Delta_z \\ \Delta_y \end{bmatrix} = \begin{bmatrix} \Delta_{1m} \\ \Delta_{1n} \\ \Delta_{2m} \\ \Delta_{2n} \end{bmatrix} \quad (2.5)$$

266 where  $\Delta_{cm}$  and  $\Delta_{cn}$  are the two displacement components estimated from images from  
 267 the  $c$  camera ( $c = 1, 2$ ),  $\Delta_x$ ,  $\Delta_y$  and  $\Delta_z$  (pixels) are the three displacement components  
 268 in the  $x$ ,  $y$ ,  $z$  directions (figure 1a) and  $\psi_{cm}$  and  $\psi_{cn}$  are calibration factors calculated  
 269 at the centre of each interrogation region using the camera calibration model. The  $\psi_{cm}$   
 270 and  $\psi_{cn}$  values indicate the shift in dewarped image coordinate (respectively in the  $m$   
 271 and  $n$  directions) corresponding to a unit displacement in the  $y$  (out-of-plane) direction  
 272 and are equivalent to the tangents of the local view angles.

273 Equation (2.5) is an overdetermined system of linear equations (4 equations, 3 un-  
 274 knowns). It can be solved using a least squares method (Raffel *et al.* 2007) or by calcu-  
 275 lating exact solutions to subsets of the four equations (e.g. Prasad 2000). In the latter  
 276 case, a redundant estimate for one of the velocity components may be obtained and it  
 277 is standard practice to average together the redundant estimates to reduce the variance  
 278 of the measurement noise in that component by a factor of two (Prasad 2000). More  
 279 efficient use of the redundancy in (2.5) can be made by storing the redundant estimates  
 280 separately rather than averaging them together. Following the method introduced for  
 281 acoustic Doppler velocimeters by Hurther & Lemmin (2001), some velocity statistics can  
 282 then be calculated which have significantly reduced noise contribution. This approach has  
 283 not previously been tested with stereoscopic PIV data and so a brief evaluation is given  
 284 in the following section. For the present camera configuration, the redundancy inherent  
 285 in (2.5) falls substantially on the  $\Delta_x$  displacement component, although in the general



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286 case, it may be shared between all of the displacement components. Based on (2.5), we  
287 can write:

$$288 \quad v = \Delta_y (M \Delta_{ls})^{-1} = (\Delta_{1n} - \Delta_{2n}) (\psi_{1n} - \psi_{2n})^{-1} (M \Delta_{ls})^{-1} \quad (2.6)$$

$$289 \quad w = \Delta_z (M \Delta_{ls})^{-1} = (\Delta_{1n}) (M \Delta_{ls})^{-1} - \psi_{1n} v = (\Delta_{2n}) (M \Delta_{ls})^{-1} - \psi_{2n} v \quad (2.7)$$

$$290 \quad u_{[1]} = (\Delta_{1m}) (M \Delta_{ls})^{-1} - \psi_{1m} v \quad (2.8)$$

$$291 \quad u_{[2]} = (\Delta_{2m}) (M \Delta_{ls})^{-1} - \psi_{2m} v \quad (2.9)$$

$$292 \quad u = 0.5 (u_{[1]} + u_{[2]}) \quad (2.10)$$

292 where  $M$  is a scale factor of the dewarped images (pixels/mm),  $\Delta_{ls}$ (ms) is the time  
293 separation between laser pulses,  $u_{[1]}$  and  $u_{[2]}$  are redundant estimates of the  $u$  velocity  
294 component, and  $u$ ,  $v$ , and  $w$  are the velocity components (m/s) in the  $x$ ,  $y$ , and  $z$   
295 directions respectively.

### 296 2.5. Noise reduction

297 The redundancy in the streamwise velocity measurement can be used to calculate velocity  
298 variance with a substantially reduced contribution of measurement noise. The instanta-  
299 neous measured velocity fluctuation ( $u' = u - \bar{u}$ ) can be decomposed into the sum of the  
300 actual velocity fluctuation ( $u_a'$ ) and the measurement error ( $\varepsilon_u'$ ) as  $u' = u_a' + \varepsilon_u'$ . The  
301 measured velocity variance can then be written:

$$\overline{u'u'} = \overline{(u_a' + \varepsilon_u')(u_a' + \varepsilon_u')} = \overline{u_a'u_a'} + \overline{\varepsilon_u'\varepsilon_u'} + 2\overline{u_a'\varepsilon_u'} \quad (2.11)$$

302 where the term  $2\overline{u_a'\varepsilon_u'}$  vanishes if the measurement error is not correlated with the ac-  
303 tual velocity fluctuation. The measured velocity variance therefore includes contributions  
304 from the actual velocity variance and the variance of the random measurement error. If  
305 redundant estimates of the velocity fluctuation ( $u_{[1]}', u_{[2]}'$ ) are available, (2.11) can be  
306 rewritten as:

$$\overline{u_{[1]}'u_{[2]}'} = \overline{(u_a' + \varepsilon_{u_{[1]}}')(u_a' + \varepsilon_{u_{[2]}}')} = \overline{u_a'u_a'} + \overline{\varepsilon_{u_{[1]}}'\varepsilon_{u_{[2]}}'} + \overline{u_a'\varepsilon_{u_{[1]}}'} + \overline{u_a'\varepsilon_{u_{[2]}}'} \quad (2.12)$$

307 where  $\varepsilon_{u_{[1]}}'$  and  $\varepsilon_{u_{[2]}}'$  are the measurement errors associated with  $u_{[1]}'$  and  $u_{[2]}'$  respec-  
308 tively. The third and fourth terms on the right vanish if the measurement error is not  
309 correlated with the velocity fluctuation, leaving  $\overline{\varepsilon_{u_{[1]}}'\varepsilon_{u_{[2]}}'}$  as the noise contribution to  
310 the measured velocity variance. The magnitude of  $\overline{\varepsilon_{u_{[1]}}'\varepsilon_{u_{[2]}}'}$  depends on the degree of  
311 correlation ( $-1 \leq C_{\varepsilon_{u_{[1]2}}} \leq 1$ ) between the two noise terms, i.e.:

$$\overline{\varepsilon_{u_{[1]}}'\varepsilon_{u_{[2]}}'} = C_{\varepsilon_{u_{[1]2}}} \sqrt{(\overline{\varepsilon_{u_{[1]}}'\varepsilon_{u_{[1]}}'}) (\overline{\varepsilon_{u_{[2]}}'\varepsilon_{u_{[2]}}'})} \quad (2.13)$$

312 In the ideal case  $C_{\varepsilon_{u_{[1]2}}}$  approaches zero, and if additionally the measurement error is  
313 uncorrelated with the velocity fluctuation, then  $\overline{u_{[1]}'u_{[2]}'}$  can be considered a ‘noise free’  
314 estimate of the velocity variance. In practice, although  $u_{[1]}$  and  $u_{[2]}$  are measured by  
315 different cameras, some correlation between the noise terms might be expected as the  
316 same particles are imaged by both cameras, albeit from different angles. Furthermore  
317 the equations for  $u_{[1]}$  and  $u_{[2]}$  (2.8, 2.9) both include the transverse velocity  $v$ . In the  
318 present study, however, the multiplying factors  $\psi_{1m}$  and  $\psi_{2m}$  are quite small, increasing  
319 from zero at the centre of the image to around  $|\psi_{cm}| = 0.2$  at the left and right edges.  
320 Nevertheless, part of the measurement error in  $v$  will contribute to  $\overline{u_{[1]}'u_{[2]}'}$ .

321 For the camera configuration used in our field experiments, this approach is limited  
322 to reducing the noise in statistics of the streamwise velocity component. The noise level

323 in the other components can, however, still be estimated. Based on (2.5), by assuming  
 324  $\psi_{1n} = -\psi_{2n}$  (for a symmetric camera system) and  $\psi_{1m} = \psi_{2m} = 0$ , and applying  
 325 standard equations for error propagation, it can be shown that

$$\overline{\varepsilon_w' \varepsilon_w'} = \psi_{1n}^2 \overline{\varepsilon_v' \varepsilon_v'} = 0.5 N_{nm} \overline{\varepsilon_{u_{[1]}' \varepsilon_{u_{[1]}'}} = 0.5 N_{nm} \overline{\varepsilon_{u_{[2]}' \varepsilon_{u_{[2]}'}} \quad (2.14)$$

326 and

$$\overline{\varepsilon_v' \varepsilon_w'} = \overline{\varepsilon_u' \varepsilon_w'} = \overline{\varepsilon_u' \varepsilon_v'} \sim 0 \quad (2.15)$$

327 where  $N_{nm} = \overline{\varepsilon_{\Delta_{cn}'} \varepsilon_{\Delta_{cn}'}} / \overline{\varepsilon_{\Delta_{cm}'} \varepsilon_{\Delta_{cm}'}}$ ,  $\varepsilon_{\Delta_{cm}'}$  is the random error in  $\Delta_{cm}$ , and  $\varepsilon_{\Delta_{cn}'}$   
 328 is the random error in  $\Delta_{cn}$ , and it is assumed that the error variance is the same for each  
 329 camera (i.e.  $\overline{\varepsilon_{\Delta_{1n}'} \varepsilon_{\Delta_{1n}'}} = \overline{\varepsilon_{\Delta_{2n}'} \varepsilon_{\Delta_{2n}'}}$  and  $\overline{\varepsilon_{\Delta_{1m}'} \varepsilon_{\Delta_{1m}'}} = \overline{\varepsilon_{\Delta_{2m}'} \varepsilon_{\Delta_{2m}'}}$ ). For the present  
 330 camera configuration, the value of  $N_{nm}$  is likely to be greater than one due to the  
 331 elongation of particle images induced by the image dewarping process. Its value can be  
 332 estimated with the help of computer generated PIV images.

333 Artificial PIV images were generated using a procedure described in Cameron (2011),  
 334 but extended here to generate a stereoscopic pair of images by applying the camera  
 335 calibration model to transform simulated three-dimensional particle coordinates to image  
 336 coordinates for a pair of cameras. The simulated cameras were positioned similar to the  
 337 real cameras used in the field experiments (63 degree viewing angle). Other parameters of  
 338 the simulation were: seeding concentration of  $9 \times 10^{-3}$  particles per pixel, particle image  
 339 diameter of 2.1 pixels, background intensity of 6 grey levels (8 bit quantization), random  
 340 additive noise with standard deviation 1.4 grey levels, maximum particle brightness of  
 341 500 grey levels (reflecting some saturation of the 8 bit image), and fill factor of 0.6. These  
 342 parameters were selected to approximate the experimental PIV images obtained in the  
 343 field. A series of  $256 \times 256$  pixel images were generated, each with a uniform displacement  
 344 field across the image, but with the displacement systematically varied over a set of  
 345  $4 \times 10^5$  images to uniformly cover the range  $0 < \Delta_{cm} < 2$  and  $0 < \Delta_{cn} < 4.4$  pixels  
 346 which corresponds to two full cycles of the peak locking error (Raffel *et al.* 2007) in each  
 347 direction. Note that the peak locking error typically has a period of 1 pixel, but when  
 348 images are dewarped, by in this case stretching the image by a factor of 2.2 in the  $n$   
 349 direction, the peak locking period is stretched by the same factor. The simulated images  
 350 were analysed using the same algorithm as was used for the field experiment images, and  
 351 the error variance for each component was obtained as:

$$\overline{\varepsilon_{u_{[1]}' \varepsilon_{u_{[1]}'}} = \overline{\varepsilon_{u_{[2]}' \varepsilon_{u_{[2]}'}} = 1.06 \overline{\varepsilon_v' \varepsilon_v'} = 0.27 \overline{\varepsilon_w' \varepsilon_w'} \quad (2.16)$$

352 indicating a value of  $N_{nm} = 7.4$ . This relationship is used in section 3.1 to estimate  
 353 the variance of the errors in the vertical and transverse velocity components. From the  
 354 simulation, the correlation coefficient between the errors in  $u_{[1]}$  and  $u_{[2]}$  was found to  
 355 be very small ( $C_{\varepsilon_{u_{[12]}'}} = 1.4 \times 10^{-3}$ ). The simulation data also indicates that the ratio  
 356  $\overline{u_a' \varepsilon_{u_{[1]}'}} / \overline{u_a' u_a'} \approx \overline{u_a' \varepsilon_{u_{[2]}'}} / \overline{u_a' u_a'}$  is of the order  $10^{-6}$  for the present experiments, con-  
 357 firming that the third and fourth terms on the right hand side of (2.12) can be safely  
 358 neglected. It is therefore reasonable to assume that  $\overline{u_{[1]}' u_{[2]}'}$  has significantly reduced  
 359 noise contribution compared to  $\overline{u_{[1]}' u_{[1]}'}$  and  $\overline{u_{[2]}' u_{[2]}'}$

360 Higher order statistics can also be estimated using redundant velocity estimates to  
 361 reduce noise contribution. For example, by calculating the velocity skewness ( $S$ ) and  
 362 kurtosis ( $K$ ) as:

$$S_u = \frac{\overline{u' u' u'}}{(\overline{u_{[1]}' u_{[2]}'})^{3/2}} \quad (2.17)$$

363

$$K_u = \frac{\overline{u_{[1]}'u_{[1]}'u_{[2]}'u_{[2]}'}} + (\overline{u_{[1]}'u_{[2]}'}) (\overline{u_{[1]}'u_{[2]}'}) - (\overline{u_{[1]}'u_{[1]}'}) (\overline{u_{[2]}'u_{[2]}'})}{(\overline{u_{[1]}'u_{[2]}'})^2} - 3 \quad (2.18)$$

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The measurement noise is eliminated if it is uncorrelated with the velocity fluctuation and if the noise correlation  $C_{\varepsilon_{u_{[12]}}}$  is zero. Again, (2.17) and (2.18) can only be applied for the streamwise velocity component for the present camera configuration. The noise contribution to the measured skewness and kurtosis for other velocity components can be estimated assuming that the random errors have a Gaussian distribution, using:

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$$S_{va} = \frac{\overline{v_a'v_a'v_a'}}{v_a'v_a'^{3/2}} = (1 + N_v)^{3/2} S_v \quad (2.19)$$

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$$S_{wa} = \frac{\overline{w_a'w_a'w_a'}}{w_a'w_a'^{3/2}} = (1 + N_w)^{3/2} S_w \quad (2.20)$$

371

$$K_{va} = \frac{\overline{v_a'v_a'v_a'v_a'}}{v_a'v_a'^2} - 3 = (1 + N_v)^2 K_v \quad (2.21)$$

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$$K_{wa} = \frac{\overline{w_a'w_a'w_a'w_a'}}{w_a'w_a'^2} - 3 = (1 + N_w)^2 K_w \quad (2.22)$$

where  $S_v$  and  $S_w$  are the measured transverse and vertical velocity skewness,  $K_v$  and  $K_w$  are the measured transverse and vertical velocity kurtosis. Actual (or noise free) velocity fluctuations ( $v_a'$ ,  $w_a'$ ), skewness ( $S_{va}$ ,  $S_{wa}$ ) and kurtosis ( $K_{va}$ ,  $K_{wa}$ ) are denoted with the subscript 'a'. The terms  $N_v$  and  $N_w$  are the noise to signal ratios defined as:

$$N_v = \frac{\overline{\varepsilon_v'\varepsilon_v'}}{v_a'v_a'} \quad \text{and} \quad N_w = \frac{\overline{\varepsilon_w'\varepsilon_w'}}{w_a'w_a'} \quad (2.23)$$

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which can be estimated from (2.16) and the experimentally measured velocity variance. These relationships are used in section 4.3 to estimate noise contributions to measured velocity skewness and kurtosis.

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### 3. Field site and experiments

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The site selected for the field deployment was on the Urie River, near the town of Inverurie and 26 km from Aberdeen City. An approximately straight section of the River was identified (figure 3a) with convenient vehicle access and a rich abundance of aquatic plants, including species from the *Myriophyllum*, *Ranunculus*, *Potamogeton*, and *Callitriche* genera, along with various aquatic mosses (figure 3b). The gravel bed at this River reach had a median particle size of 35 mm (estimated from a random sample of 117 particles) and featured intermittent sandy patches and occasional large boulders. At the test section the River was 12.9 m wide (figure 4), the average flow depth was 0.39 m, the flow rate was 2.7 m<sup>3</sup>/s, and the water surface slope was  $1.5 \pm 0.4 \times 10^{-3}$ . The Reynolds number based on flow depth and mean velocity was  $1.52 \times 10^5$  and the Froude number was 0.28. Assessment of velocity time series (not shown) and observations of the river water surface elevation throughout the deployment suggest that the flow conditions were steady.

A set of PIV measurements were made of the flow field around a *Ranunculus penicillatus* plant patch located 2.8 m from the right river bank. The maximum dimensions of the patch were approximately 400 mm long, 200 mm wide, and 100 mm high. This particular

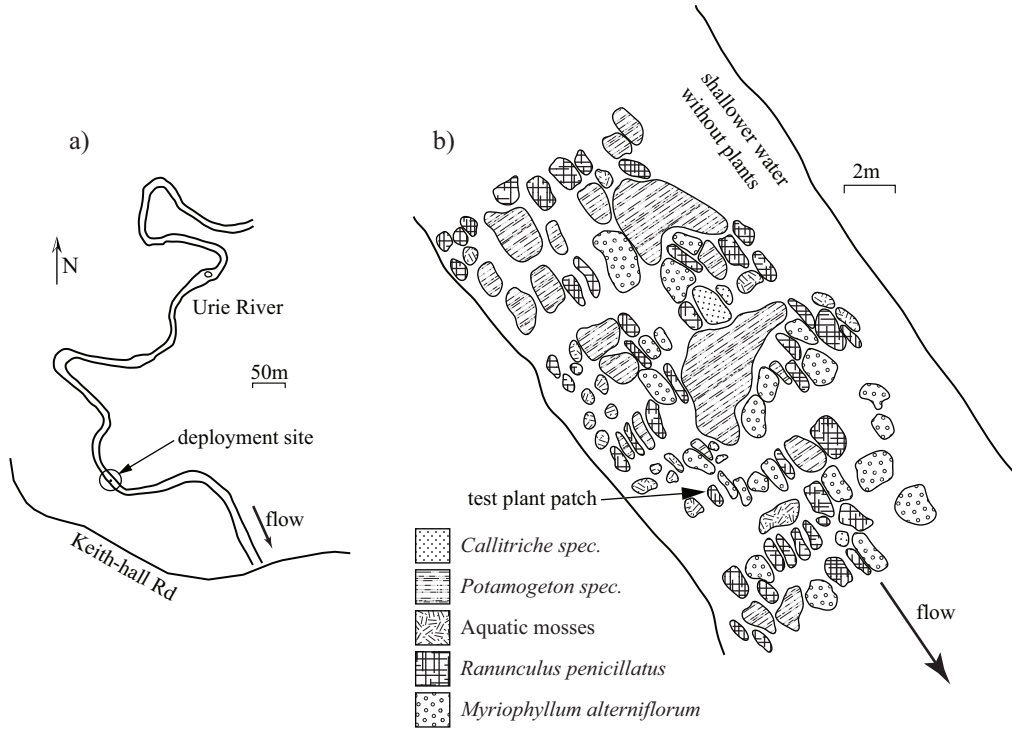


FIGURE 3. Sketch of the field deployment site (a). Aquatic plant species near the test section of the Urie River (b).

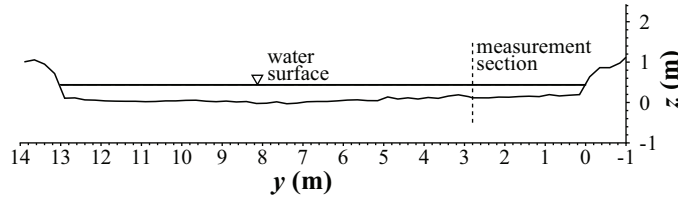


FIGURE 4. Cross section of the Urie River near the test section.

396 patch was selected because of its size in relation to the PIV field of view and its isolation  
 397 from other plants and large boulders. Three-minute PIV recordings were made at three  
 398 measurement locations, starting near the free end of the plant patch and subsequently  
 399 incremented by 130 mm in the downstream direction. The total measurement coverage  
 400 was a planar region 400 mm in the streamwise direction and 320 mm in the vertical  
 401 direction and aligned with the centreline of the plant (figure 5a). The recording rate was  
 402 30 image pairs per second, but due to a technical issue, some of the frames were later  
 403 found to not be viable, resulting in an average of 20 image pairs per second. Missing time  
 404 steps are assigned  $\phi_M = 0$  allowing statistical quantities to be estimated using only valid  
 405 data.

### 406 3.1. Measurement noise

407 Based on the redundant estimates of the streamwise velocity component, the variance of  
 408 the noise can be estimated as:

$$0.5 (\overline{\varepsilon_{u_{[1]}' \varepsilon_{u_{[1]}'}} + \varepsilon_{u_{[2]}' \varepsilon_{u_{[2]}'}}}) = 0.5 (\overline{u_{[1]}' u_{[1]}' + u_{[2]}' u_{[2]}' - 2u_{[1]}' u_{[2]}'}) \quad (3.1)$$

12

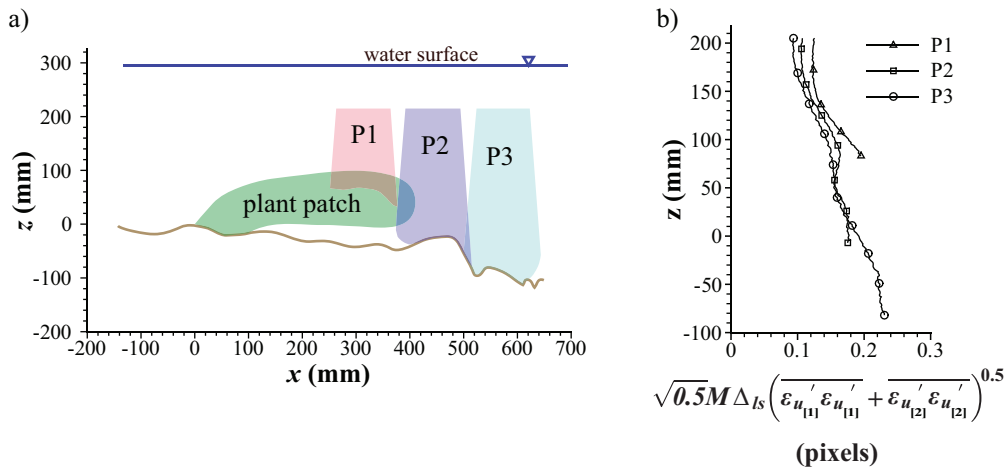


FIGURE 5. PIV measurement coverage (P1, P2, P3) relative to the plant patch and local bed topography (a). Standard deviation of measurement noise for the three measurement positions (b).

409 This noise term is plotted in the form  $\sqrt{0.5M\Delta_{l_s}(\overline{\varepsilon_{u_{[1]}'\varepsilon_{u_{[1]}'}} + \overline{\varepsilon_{u_{[2]}'\varepsilon_{u_{[2]}'}}})^{0.5}}$  in figure 5b,  
 410 which is the standard deviation of the error in displacement units (pixels), allowing compar-  
 411 ison with previous studies of PIV error (here  $M = 12$  pixels/mm and  $\Delta_{l_s} = 1$  ms).  
 412 Figure 5b indicates that the standard deviation of the measurement noise in the stream-  
 413 wise displacement component is approximately the same for each of the three measure-  
 414 ment positions and increases from around 0.1 pixels for large  $z$  (near the free surface)  
 415 to around 0.2 pixels near the bed. The increase in error approaching the bed reflects  
 416 the varying magnification of the source images and deteriorating image conditions with  
 417 distance from the cameras due to light sheet intensity falloff. The magnitude of the error  
 418 is comparable to that obtained from computer simulations when considering a significant  
 419 out of plane displacement component (e.g. Nobach & Bodenschatz 2009; Cameron 2011).  
 420 Values of the noise to signal ratio terms (2.23) for the transverse and vertical velocity  
 421 components can be estimated as  $N_v = 0.02$  and  $N_w = 0.07$  in the wake of the plant  
 422 patch.

## 423 4. Flow turbulence and plant fluctuations

### 424 4.1. Mean velocity field

425 Mean velocity streamlines combined for the three measurement planes (figure 6) indicate  
 426 that the flow does not separate from the plant patch and no recirculation zone forms.  
 427 Folkard (2011a) defines this as the ‘canopy through-flow’ regime, but the flow and patch  
 428 conditions for its existence are yet to be identified for real plants. In contrast, the small  
 429 rock immediately behind the plant patch shows clear signs of separation and recircula-  
 430 tion, highlighting the potentially different mechanisms of drag for these two objects.  
 431 Bluff bodies, such as the rock behind the plant patch, produce drag mainly through the  
 432 differential pressure between their upstream and downstream surfaces which occurs due  
 433 to flow separation. Drag on aquatic plants, however, due to their flexibility, porosity, and  
 434 large wetted surface area, may be dominated by viscous drag (Nikora & Nikora 2007)  
 435 which forms due to the velocity gradient at the plant surfaces. Although figure 6 is con-  
 436 sistent with the proposed conjecture, this hypothesis is difficult to test experimentally  
 437 as flow separation and pressure drag may occur at several different plant scales (plant

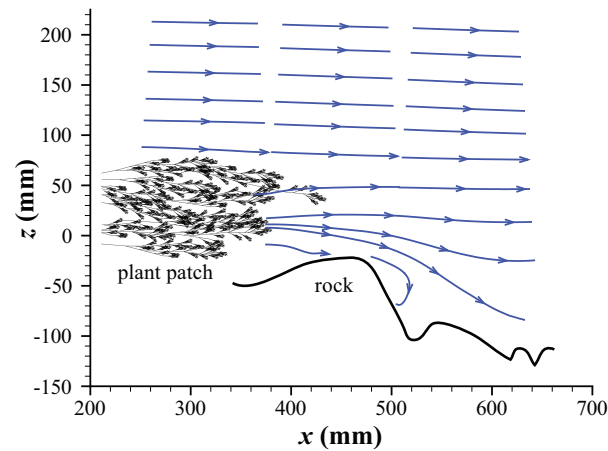


FIGURE 6. Velocity streamlines in the flow region around a *Ranunculus* plant patch.

438 patch, individual plant, stem, leaf). Recent experimental studies have measured drag  
 439 forces at each of these scales (e.g. Albayrak *et al.* 2012; Nikora *et al.* 2012; Siniscalchi &  
 440 Nikora 2012; Siniscalchi *et al.* 2012; Siniscalchi & Nikora 2013), however, separating vis-  
 441 cous/pressure drag contributions directly still exceeds experimental capability. Further  
 442 complicating the viscous/pressure drag argument is that simple scaling relationships such  
 443 as  $F_D \propto u^2$  for pressure drag and  $F_D \propto u^1$  for viscous drag (where  $F_D$  is drag force on  
 444 the plant) cannot easily be applied to aquatic plants as they have the tendency to change  
 445 their structure in response to the velocity field. This so called ‘reconfiguration’ (Vogel  
 446 1994; de Langre 2008) can change the wetted surface area, the effective frontal area and  
 447 the drag coefficient (through streamlining) of the plant as a function of flow velocity,  
 448 thereby complicating interpretation of force scaling with flow velocity.

449 The plant drag, whether viscous or pressure dominated, is a sink of momentum and  
 450 introduces a free shear layer (and associated inflection in the  $\bar{u}(z)$  profile) at the interface  
 451 between the retarded flow in the wake of the plant and the background channel flow  
 452 (figure 7). The inflectional form of the mean velocity profile is suggested to lead to  
 453 the Monami phenomenon in aquatic plant canopies (Ghisalberti & Nepf 2002; Nezu &  
 454 Sanjou 2008; Nepf 2012) and dominate local turbulence characteristics due to a periodic  
 455 production of vortices (Kelvin Helmholtz instability). For a single isolated plant patch,  
 456 however, the mean flow in the wake is distinctly three-dimensional and exposed to high  
 457 background turbulence levels which would tend to disrupt any periodic vortex formation  
 458 mechanisms. The shear layer may nevertheless be associated with high levels of turbulence  
 459 production; the distribution of the Reynolds stresses and their budget terms are examined  
 460 in the following section. Potential periodicity of the velocity in the plant wake is examined  
 461 in section 4.5. Figure 7b illustrates the streamwise momentum recovery in the wake of  
 462 the plant for the  $x$  coordinates marked by circles in figure 7a. The streamwise velocity  
 463 in the wake is steadily increasing with increasing  $x$ , and the corresponding decay of the  
 464 maximum velocity gradient is apparent.

#### 4.2. Reynolds stresses and their budget terms

465 The normal Reynolds stresses ( $\overline{u'u'}$ ,  $\overline{v'v'}$ ,  $\overline{w'w'}$ ) and the primary Reynolds shear stress  
 466 ( $-\overline{u'w'}$ ) all attain maximum values near the shear layer in the wake of the plant (figure  
 467 8). In general, the distribution of Reynolds stresses in the patch wake may depend on  
 468 a variety of patch and approach flow conditions such as patch length and width, the  
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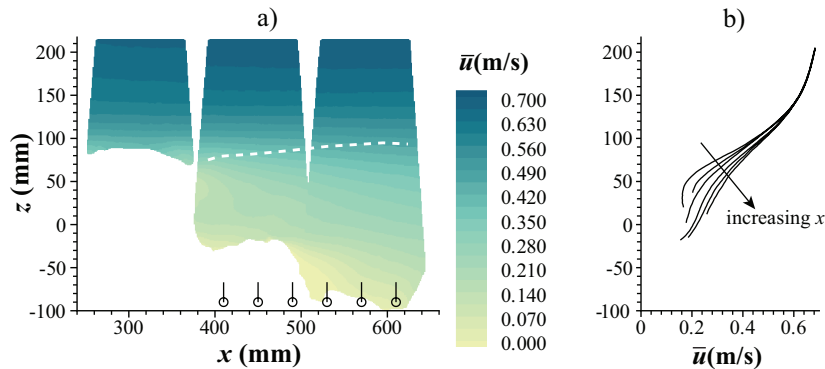


FIGURE 7. Time-averaged streamwise velocity distribution: a) around a *Ranunculus* plant patch, and b) in the wake of the plant patch for  $x$  coordinates corresponding to circle symbols in ‘a’. Dashed line indicates local maximum in the  $\partial\bar{u}/\partial z$  distribution.

470 distribution and shape of plant stems and leaves within the patch, the flexibility of  
 471 the plants, the approach flow Reynolds number, and the flow depth to patch height  
 472 ratio. For example, in contrast to our study, the peak Reynolds stress for Folkard’s  
 473 (2005) model seagrass canopy formed several patch heights downstream of the patch  
 474 and near the reattachment point of the separated flow. It is not yet clear if natural  
 475 patches of *Ranunculus penicillatus* form similar wake features under different flow and  
 476 patch conditions. Secondary Reynolds shear stresses (not shown) were found to be an  
 477 order of magnitude smaller than the primary Reynolds shear stress as might be expected  
 478 (due to symmetry) near the centreline of the plant. The Reynolds stress correlation  
 479 coefficient  $-\overline{u'w'}/(\overline{u'u'}\overline{w'w'})^{0.5}$ , which reflects the efficiency of the turbulent fluctuations  
 480 at redistributing momentum, has a maximum value of 0.61 in the plant wake ( $x =$   
 481  $400, z = 75$ ), slightly larger than the 0.4-0.5 typical for open channel flows (Nezu &  
 482 Nakagawa 1993), the 0.5 found for terrestrial canopies (Raupach *et al.* 1996), and the 0.5  
 483 found in the wake of a cylinder (Cantwell & Coles 1983).

484 The transverse and vertical normal stresses have similar magnitudes to each other in  
 485 the plant wake ( $\overline{v'v'}/\overline{w'w'} = 1 - 1.2$ ) which is smaller than the ratio 1.65 typical for open  
 486 channel flows (Nezu & Nakagawa 1993), but closer to the ratio of 1.2 measured for a  
 487 plane mixing layer by Wygnanski & Fiedler (1970). The streamwise normal stress  $\overline{u'u'}$   
 488 is found to decay with increasing  $x$  much faster than the other components. In the far  
 489 wake ( $x = 600$  mm)  $\overline{u'u'}$  has reduced to 71% of its near wake ( $x = 400$  mm) maximum.  
 490 Corresponding values for  $\overline{v'v'}$  and  $\overline{w'w'}$  are 91% and 85% respectively. Subtle differences  
 491 in the elevations where the maximum variance occurs can be seen between the different  
 492 components of the Reynolds stress tensor. Local maximums in both  $\overline{u'u'}$  ( $z$ ) and  $\overline{u'w'}$  ( $z$ )  
 493 tend to higher elevations with increasing  $x$  following the mean shear layer and reflecting  
 494 the expansion of the wake region into the outer flow. The trend for  $\overline{w'w'}$  ( $z$ ) is nearly  
 495 horizontal, and for  $\overline{v'v'}$  ( $z$ ) it is downward. The reason for these different trends is not  
 496 clear, but further understanding might be gained by considering the budget equation for





509 resulting equation. Third, exchange the free indices ( $i$  and  $k$ ) in the equation developed  
 510 in step 2 and add this new equation to the original equation in step 2 to give (4.1). It can  
 511 be noted that the budget equation for turbulent kinetic energy is obtained by taking half  
 512 the trace of (4.1). Equation (4.1) has received considerable attention as a framework to  
 513 develop closure models for the Reynolds averaged Navier-Stokes equations. Distribution  
 514 of the terms in (4.1) can also provide some insight into the turbulence in the wake of the  
 515 plant patch, and in the present study this is our primary interest.

516 In a uniform, two-dimensional channel flow  $\bar{v} = \bar{w} = \partial/\partial x = \partial/\partial y = 0$  and the only  
 517 non-zero normal stress production term is in the  $\overline{u'u'}$  budget. Variance is redistributed  
 518 from  $\overline{u'u'}$  to the other normal stress components by the pressure-strain correlation term  
 519 which is traceless and therefore does not appear in the total turbulent kinetic energy  
 520 balance. Away from boundaries, the dissipation rate is expected to be approximately  
 521 equal in each of the normal stress budgets due to local isotropy, if the Reynolds number  
 522 is reasonably high (Davidson 2004). Dissipation in the  $\overline{u'u'}$  budget is typically small, and  
 523 production in  $\overline{u'u'}$  is balanced largely by the pressure-strain term (Mansour *et al.* 1988;  
 524 Pope 2000). The mean convection, turbulent transport, and pressure transport terms  
 525 act to redistribute the Reynolds stresses in space and each of the transport terms inte-  
 526 grate to zero over the flow depth in two-dimensional channel flow. The viscous transport  
 527 term is expected to be negligible away from boundaries compared to other transport  
 528 mechanisms if the Reynolds number is large. In the wake of an aquatic plant patch, the  
 529 time-averaged flow field is three-dimensional and some departure from the distributions  
 530 of the budget terms for two-dimensional flow may be expected. Some of the terms in  
 531 (4.1) cannot be evaluated from the experimental data. The pressure field is not avail-  
 532 able, terms involving transverse derivatives cannot be calculated, and there is insufficient  
 533 spatial resolution to resolve the dissipation rate tensor that would require resolution of  
 534 the order of the Kolmogorov microscale ( $\sim 0.1$  mm). We can, however, estimate con-  
 535 tributions from streamwise and vertical derivatives to the mean convection, turbulent  
 536 transport, and production terms as (4.2)-(4.5), where the terms in brackets highlight the  
 537 transverse derivatives that could not be calculated in this study. The effect of random  
 538 measurement errors should be negligible for the terms involving third moments (all tur-  
 539 bulent transport terms) and terms involving the fluid stresses  $\overline{u_{[1]}'u_{[2]}'}$ ,  $\overline{u'u'}$ ,  $\overline{u'v'}$ , or  $\overline{v'w'}$ .  
 540 Terms involving  $\overline{v'v'}$  or  $\overline{w'w'}$  will be biased by the measurement noise, but evaluation  
 541 of the magnitude of the noise contribution to each of these terms suggests that in all  
 542 cases it is much smaller than the sampling error. Sampling errors were estimated using  
 543 a resampling technique (Garcia *et al.* 2006) and associated confidence intervals are indi-  
 544 cated in figure 9. In general, the sampling error varies with  $z$ , but in order to reduce the  
 545 clutter in figure 9, an average value is given. Derivatives were estimated by convolving  
 546 the time averaged moments of the velocity field with a  $21 \times 21$  grid point ( $21 \times 21$  mm)  
 547  $2^{nd}$  order least squares kernel. The size of the filter was sufficiently large to smooth over  
 548 sampling errors (due to finite measurement duration), but still sufficiently small so as  
 549 not to significantly reduce the amplitude of the measured derivatives.

$$\begin{array}{l}
\text{mean conv.} \\
\text{turb. trans.} \\
\text{production}
\end{array}
\left. \begin{array}{l}
-\rho\bar{u}\frac{\partial\overline{u'_{[1]}u'_{[2]}}}{\partial x} - \rho\bar{w}\frac{\partial\overline{u'_{[1]}u'_{[2]}}}{\partial z} - \left(\rho\bar{v}\frac{\partial\overline{u'u'}}{\partial y}\right) \\
-\rho\frac{\partial\overline{u'u'u'}}{\partial x} - \rho\frac{\partial\overline{w'u'u'}}{\partial z} - \left(\rho\frac{\partial\overline{v'u'u'}}{\partial y}\right) \\
-2\rho\overline{u'_{[1]}u'_{[2]}}\frac{\partial\bar{u}}{\partial x} - 2\rho\overline{w'u'}\frac{\partial\bar{u}}{\partial z} - \left(2\rho\overline{u'v'}\frac{\partial\bar{u}}{\partial y}\right)
\end{array} \right\} \rho\overline{u'u'} \text{ budget} \quad (4.2)$$

$$\begin{array}{l}
\text{mean conv.} \\
\text{turb. trans.} \\
\text{production}
\end{array}
\left. \begin{array}{l}
-\rho\bar{u}\frac{\partial\overline{v'v'}}{\partial x} - \rho\bar{w}\frac{\partial\overline{v'v'}}{\partial z} - \left(\rho\bar{v}\frac{\partial\overline{v'v'}}{\partial y}\right) \\
-\rho\frac{\partial\overline{u'v'v'}}{\partial x} - \rho\frac{\partial\overline{w'v'v'}}{\partial z} - \left(\rho\frac{\partial\overline{v'v'v'}}{\partial y}\right) \\
-2\rho\overline{u'v'}\frac{\partial\bar{v}}{\partial x} - 2\rho\overline{w'v'}\frac{\partial\bar{v}}{\partial z} - \left(2\rho\overline{v'v'}\frac{\partial\bar{v}}{\partial y}\right)
\end{array} \right\} \rho\overline{v'v'} \text{ budget} \quad (4.3)$$

$$\begin{array}{l}
\text{mean conv.} \\
\text{turb. trans.} \\
\text{production}
\end{array}
\left. \begin{array}{l}
-\rho\bar{u}\frac{\partial\overline{w'w'}}{\partial x} - \rho\bar{w}\frac{\partial\overline{w'w'}}{\partial z} - \left(\rho\bar{v}\frac{\partial\overline{w'w'}}{\partial y}\right) \\
-\rho\frac{\partial\overline{u'w'w'}}{\partial x} - \rho\frac{\partial\overline{w'w'w'}}{\partial z} - \left(\rho\frac{\partial\overline{v'w'w'}}{\partial y}\right) \\
-2\rho\overline{u'w'}\frac{\partial\bar{w}}{\partial x} - 2\rho\overline{w'w'}\frac{\partial\bar{w}}{\partial z} - \left(2\rho\overline{v'w'}\frac{\partial\bar{w}}{\partial y}\right)
\end{array} \right\} \rho\overline{w'w'} \text{ budget} \quad (4.4)$$

$$\begin{array}{l}
\text{mean conv.} \\
\text{turb. trans.} \\
\text{production}
\end{array}
\left. \begin{array}{l}
-\rho\bar{u}\frac{\partial\overline{u'w'}}{\partial x} - \rho\bar{w}\frac{\partial\overline{u'w'}}{\partial z} - \left(\rho\bar{v}\frac{\partial\overline{u'w'}}{\partial y}\right) \\
-\rho\frac{\partial\overline{u'u'w'}}{\partial x} - \rho\frac{\partial\overline{w'u'w'}}{\partial z} - \left(\rho\frac{\partial\overline{u'w'v'}}{\partial y}\right) \\
-\rho\overline{u'w'}\frac{\partial\bar{u}}{\partial x} - \rho\overline{w'w'}\frac{\partial\bar{u}}{\partial z} - \rho\overline{u'_{[1]}u'_{[2]}}\frac{\partial\bar{w}}{\partial x} - \rho\overline{u'w'}\frac{\partial\bar{w}}{\partial z} \\
-\left(\rho\overline{w'v'}\frac{\partial\bar{u}}{\partial y} - \rho\overline{u'v'}\frac{\partial\bar{w}}{\partial y}\right)
\end{array} \right\} \rho\overline{u'w'} \text{ budget} \quad (4.5)$$

550 Distributions of the available Reynolds stress budget terms together with the Reynolds  
551 stresses are shown in figure 9 for the near wake ( $x = 430$  mm) and for the far wake  
552 ( $x = 600$  mm). The distribution of the production term in the  $u'u'$  budget forms a peak  
553 at  $z = 85$  mm in the near wake and  $z = 99$  mm in the far wake which closely matches the  
554 peaks in the corresponding Reynolds normal stress. The mean convection term is also a  
555 gain at this elevation but is much smaller than the production term. The ratio between  
556 the local production and the mean convection terms, 3.8 in the near wake, suggests that  
557 the  $u'u'$  field near the shear layer is dominated by local rather than upstream generation  
558 processes. This result is consistent with a plane mixing layer (Wyganski & Fiedler  
559 1970), but differs from many separated flows which feature a region where convection is  
560 the dominant gain term, for example the axisymmetric wake (Uberoi & Freymuth 1970)

561 and the wake of a surface mounted cube (Hussein & Martinuzzi 1996). The turbulent  
 562 transport term in the  $\overline{u'u'}$  budget is a loss near the shear layer and a gain at both higher  
 563 and lower elevations. The effect of this term is therefore to diffuse turbulence away  
 564 from the shear layer where it is produced. It is interesting to note that both turbulent  
 565 transport and mean convection terms cross zero above and below the shear layer at  
 566 about the same elevations ( $z = 114$  mm and  $z = 64$  mm respectively). Further, these  
 567 elevations correspond to measurement porosity values measured at  $x = 350$  mm of  $\overline{\phi_M} =$   
 568  $0.98$  and  $0.04$  respectively, i.e. near the extreme upper and lower elevations of the top of  
 569 the fluctuating plant (considering that  $\overline{\phi_M}$  in this region is dominated by the presence or  
 570 absence of plant within PIV interrogation regions). The alignment between these three  
 571 statistics may indicate that the fluctuating plant (and corresponding fluctuation of the  
 572 shear layer elevation) plays a role in regulating the distribution of  $\overline{u'u'}$  in the wake.  
 573 Production in the  $\overline{v'v'}$  and  $\overline{w'w'}$  budgets is small compared to the production in the  
 574  $\overline{u'u'}$  budget and does not appear to explain why the maximums in the three Reynolds  
 575 normal stress distributions do not coincide. The reason for this is likely contained in the  
 576 pressure-strain and pressure transport terms (which are not available from experimental  
 577 data) and also in upstream production such as in the wakes of individual plant stems and  
 578 leaves. Both the turbulent transport and mean convection terms in the transverse and  
 579 vertical normal stress budgets have similar characteristics to the corresponding terms in  
 580 the streamwise normal stress budget. The transport terms are a loss where the velocity  
 581 variance is high and a gain in both the higher and lower flow layers. Convection terms  
 582 follow the same pattern, but are smaller and have opposite sign. The  $\overline{u'u'}$  budget has  
 583 similar characteristics to the  $\overline{u'u'}$  budget, but each term has opposite sign because the  
 584 primary Reynolds shear stress is negative. We note again a correlation between the  
 585 distribution of the production term and the corresponding Reynolds stress distribution  
 586 with the local peaks in these distributions forming at the same elevation. The production  
 587 is 7.3 times larger than the convection term indicating that the primary Reynolds shear  
 588 stress distribution is dominated by local rather than upstream production. The turbulent  
 589 transport and mean convection terms have opposite signs and similar to the  $\overline{u'u'}$  budget,  
 590 each crosses zero near the same elevation.

#### 591 4.3. Higher order moments

592 Skewness  $S_i = \overline{u'_i u'_i u'_i} / \overline{u'_i u'_i}^{3/2}$  and kurtosis  $K_i = \overline{u'_i u'_i u'_i u'_i} / \overline{u'_i u'_i}^2 - 3$  (repeated  
 593 index does not imply summation) distributions provide further indication of the nature  
 594 of the turbulence in the wake of the plant patch. Equations (2.19)-(2.22) indicate that  
 595 measured skewness and kurtosis are biased towards zero by the measurement noise. The  
 596 relative error in  $S_w$  is around 10% and in  $K_w$  is around 15% in the shear zone behind the  
 597 plant patch. Relative errors for  $S_v$  and  $K_v$  are 3% and 4% respectively and for  $S_u$  and  $K_u$   
 598 the error contribution is minimised using (2.17) and (2.18). Skewness is an indicator of  
 599 the asymmetry of the velocity probability distribution, with negative skewness associated  
 600 with a left-tailed distribution (rare high magnitude velocity fluctuations tend to have a  
 601 negative sign) while positive skewness indicates a right-tailed distribution (rare high  
 602 magnitude events tend to have a positive sign). Figure 9 indicates that  $S_v$  in the wake  
 603 of the plant is near zero over much of the flow depth which is expected due to the  
 604 approximate symmetry of the time averaged flow field near the plant centreline. Skewness  
 605 of the streamwise and bed-normal velocity components have opposite signs over most  
 606 of the flow depth. A transition from an ‘ejection’ dominated upper flow region ( $S_u <$   
 607  $0$ ,  $S_w > 0$ ) to a ‘sweep’ dominated lower flow region ( $S_u > 0$ ,  $S_w < 0$ ) is evident  
 608 around  $z = 88$  mm in the near wake which corresponds to the location of the mean  
 609 shear layer. Such antisymmetric distributions of  $S_u$  and  $S_w$  are typical of mixing layers

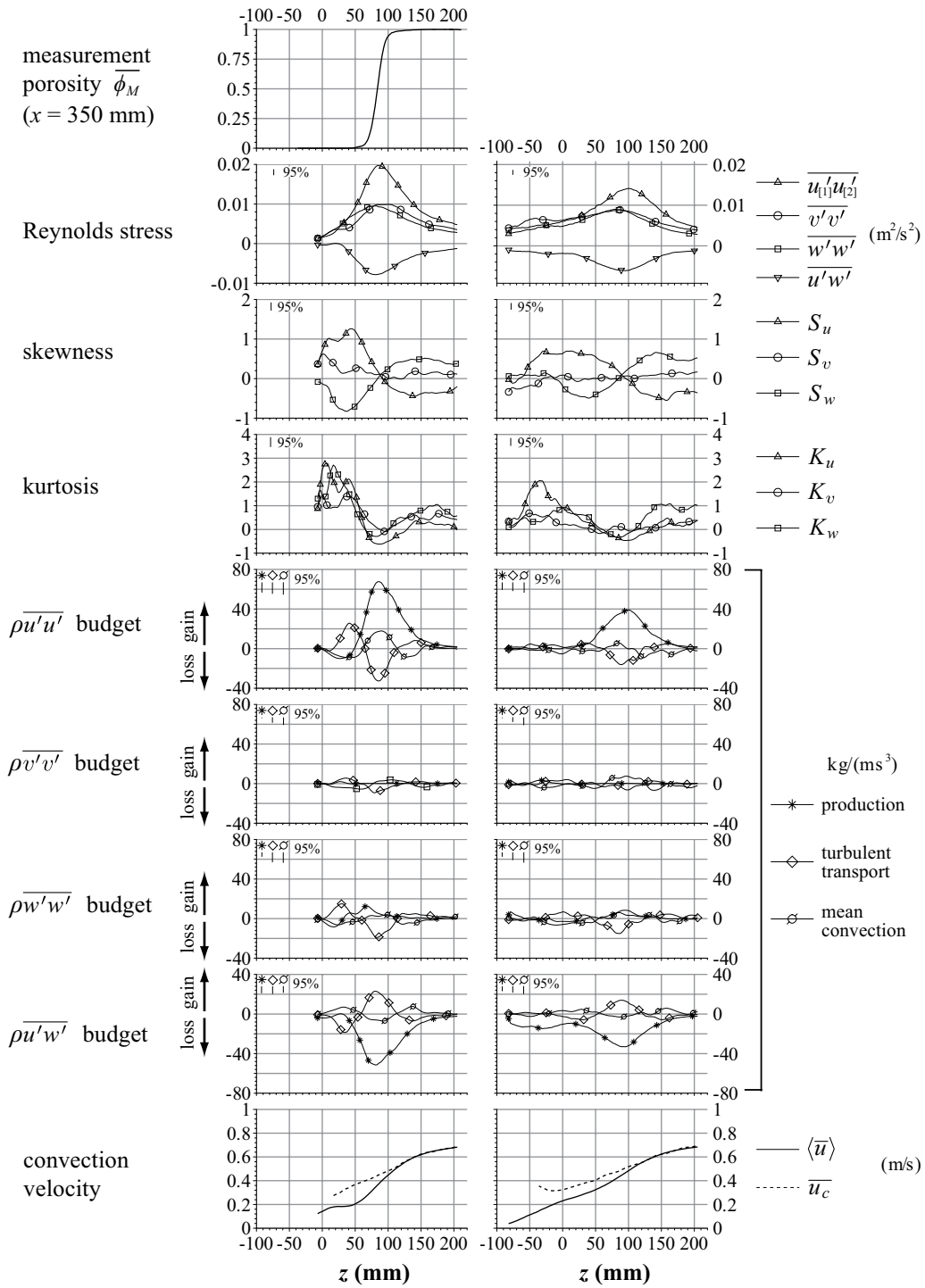


FIGURE 9. Velocity field statistics in the near wake ( $x = 430$ , left column except measurement porosity) and far wake ( $x = 600$ , right column) of a *Ranunculus penicillatus* plant patch.

20

610 and flows over aquatic canopies (Raupach *et al.* 1996; Nezu & Sanjou 2008) but are also  
 611 found in open channel flows over gravel beds (Nikora & Goring 2000*b*). The kurtosis  
 612 of a probability distribution is an indicator of its ‘peakedness’ relative to a Gaussian  
 613 distribution. A high value of the kurtosis coefficient of a velocity signal indicates the  
 614 presence of rare (intermittent) high magnitude events, while a kurtosis coefficient that  
 615 is less than zero indicates that high magnitude events occur more frequently than for a  
 616 Gaussian distribution. Figure 9 indicates that the kurtosis coefficient for each velocity  
 617 component follows a similar distribution with regions of positive kurtosis in the higher  
 618 and lower flow layers and a region of low kurtosis near the shear layer. The largest values  
 619 of kurtosis are found behind the plant for elevations between  $z = 0$  mm and  $z = 50$  mm.  
 620 This indicates, in conjunction with  $S_u > 0$ ,  $S_w < 0$ , that the flow field in this region  
 621 is characterised by rare high magnitude sweep events that likely originate from higher  
 622 flow layers and intermittently impinge into the low velocity region behind the plant. The  
 623 negative value of kurtosis near the mean shear layer ( $K_u = -0.61$  at  $z = 88$  mm) is  
 624 similar to the value of  $-0.63$  measured by Wygnanski & Fiedler (1970) at the centre of  
 625 a plane mixing layer. Negative values of kurtosis have also been found in the near-bed  
 626 region of gravel bed open channel flows by Nikora & Goring (2000*b*) and for a smooth  
 627 wall boundary layer by Balachandar *et al.* (2001). It is interesting to note that in each  
 628 of these examples, and also for the present aquatic plant wake, the location of minimum  
 629 kurtosis corresponds to the location of maximum variance.

#### 630 4.4. Convection Velocity

631 Eddy convection velocity ( $\overline{u_c}$ ) has previously been studied primarily because of its rel-  
 632 evance to Taylors ‘frozen turbulence’ approximation which can be applied to transform  
 633 velocity statistics (such as velocity spectra and correlation functions) between time and  
 634 space domains. Several studies have indicated surprising departures of the convection  
 635 velocity from the local mean velocity (with  $\overline{u_c} > \overline{u}$ ) such as in terrestrial canopy flows  
 636 by Shaw *et al.* (1995), in aquatic canopies by Nezu & Sanjou (2008), and for gravel bed  
 637 open channel flows by Nikora & Goring (2000*a*). Understanding the reasons for this de-  
 638 parture may provide some further insight into the turbulence structure, and this is our  
 639 motivation for examining the convection velocity in the wake of the *Ranunculus* plant  
 640 patch.

641 Convection velocity in the wake of the plant patch can be estimated from the 2-point  
 642 space-time correlation:

$$643 \quad R(z, x_u, x_d, \Delta t_n) \left[ \overline{u'u'}(x_u, z) \overline{u'u'}(x_d, z) \right]^{0.5} =$$

$$644 \quad \frac{\sum_{t_n=1}^{t_n=T} [u'(x_u, z, t_n)u'(x_d, z, t_n + \Delta t_n)\phi_M(x_u, z, t_n)\phi_M(x_d, z, t_n + \Delta t_n)]}{\sum_{t_n=1}^{t_n=T} [\phi_M(x_u, z, t_n)\phi_M(x_d, z, t_n + \Delta t_n)]} \quad (4.6)$$

643 where  $x_u$  and  $x_d$  identify ‘upstream’ and ‘downstream’  $x$  coordinates,  $t_n$  is the time  
 644 step,  $\Delta t_n$  is time step separation,  $T$  is the total number of time steps, and  $\phi_M$  is the  
 645 measurement clipping function described in section 2.2. The eddy convection velocity is  
 646 then:

$$647 \quad \overline{u_c}(z, x_u, x_d) = \frac{x_d - x_u}{\Delta t_{R_{\max}}} f_s \quad (4.7)$$

647 where  $\Delta t_{R_{\max}}$  is the time separation (measured by time steps) that maximises  $R$ , and  
 648  $f_s$  is the sampling frequency (30 Hz). The mean velocity field in the wake of the plant  
 649 is not homogeneous, so in order to make a meaningful comparison between the convec-  
 650 tion velocity and the local mean velocity, the latter is spatially averaged over the range

651  $x_u < x < x_d$ :

$$\langle \bar{u} \rangle (z, x_u, x_d) = \frac{1}{\sum_{x=x_u}^{x=x_d} \sum_{t_n=1}^{t_n=T} [\phi_M(x, z, t_n)]} \sum_{x=x_u}^{x=x_d} \sum_{t_n=1}^{t_n=T} [u(x, z, t_n) \phi_M(x, z, t_n)] \quad (4.8)$$

652 Convection velocity and local average velocity are shown for the near wake ( $x_u =$   
 653  $410, x_d = 450$  mm) and for the far wake ( $x_u = 580, x_d = 620$  mm) in figure 9. Below  
 654  $z = 115$  mm in the near wake and below  $z = 120$  mm in the far wake, the convection  
 655 velocity deviates significantly from the local mean velocity. The result  $\bar{u}_c > \langle \bar{u} \rangle$  might  
 656 be expected in the lower flow layers ( $z < 50$  mm) as the turbulence in this region is  
 657 characterised by rare high magnitude velocity fluctuations which are generated near the  
 658 shear layer (where the mean velocity is higher) and periodically impinge into the low  
 659 velocity region. It is reasonable to assume that these eddies propagate with a velocity  
 660 close to the mean velocity where they are generated explaining the observed  $\bar{u}_c > \langle \bar{u} \rangle$   
 661 near the bed. We note, however, that near the shear layer ( $z = 85$  mm in the near wake)  
 662 where the velocity fluctuations are dominated by local production, we can still observe  
 663 that the convection velocity is larger than the local mean velocity ( $\bar{u}_c / \langle \bar{u} \rangle = 1.2$  at  
 664  $z = 85$  mm). Raupach *et al.* (1996) explains similar observations in terrestrial canopy  
 665 flows by suggesting that eddies which dominate the two-point correlation  $R$  are produced  
 666 mainly during wind gusts and therefore naturally propagate with the higher velocity of  
 667 the gust rather than the lower mean velocity. The relevance of this interaction mechanism  
 668 between the outer flow and the shear layer eddies to the present experiment, where scale  
 669 separation is much smaller, is not clear and remains to be clarified in future experiments.

#### 670 4.5. Velocity Spectra

671 The structure of the velocity field in the wake of the plant is further examined by consid-  
 672 ering the power spectrum of velocity fluctuations,  $F_{ii}(f)$ . The spectrum can be evaluated  
 673 for velocity data with missing samples using the Lomb-Scargle method (Lomb 1976; Scargle  
 674 1982) which can be written as:

$$2F_{ii}(\omega) = \frac{\left\{ \sum_{t_n=1}^{t_n=T} u_i \phi_M \cos[\omega(t - \tau)] \right\}^2}{\sum_{t_n=1}^{t_n=T} \phi_M \cos^2[\omega(t - \tau)]} + \frac{\left\{ \sum_{t_n=1}^{t_n=T} u_i \phi_M \sin[\omega(t - \tau)] \right\}^2}{\sum_{t_n=1}^{t_n=T} \phi_M \sin^2[\omega(t - \tau)]} \quad (4.9)$$

675 where,  $t = t_n / f_s$  is the time corresponding to the  $t_n^{\text{th}}$  measurement sample,  $\omega = 2\pi f$  is  
 676 the angular frequency,  $f$  is the linear frequency, and  $\tau$  is a time lag adopted by (Scargle  
 677 1982) to enforce invariance of the spectrum to time translation of the data and simplify  
 678 the statistical behaviour, with:

$$\tan(2\omega\tau) = \frac{\sum_{t_n=1}^{t_n=T} \phi_M \sin[2\omega t]}{\sum_{t_n=1}^{t_n=T} \phi_M \cos[2\omega t]} \quad (4.10)$$

679 The Lomb-Scargle method is equivalent to estimating the spectrum by a least squares  
 680 fit of sine waves to the data and for regularly spaced data reduces to the conventional  
 681 Fourier spectrum (Scargle 1982).

682 Comparison of the velocity power spectrum near the shear layer in the wake of the  
 683 plant ( $x = 430, z = 85$  mm, figure 10a) with the spectrum at a higher elevation ( $x = 430,$   
 684  $z = 200$  mm, figure 10b), where the influence of the plant is reduced, indicates a broad  
 685 increase in energy across all resolved frequencies in the plant wake. Some flattening of  
 686 the spectrum is evident at higher frequencies due to the contribution of aliasing and  
 687 measurement noise. A subtle clustering of energy around  $f = 1$  Hz can be seen in the

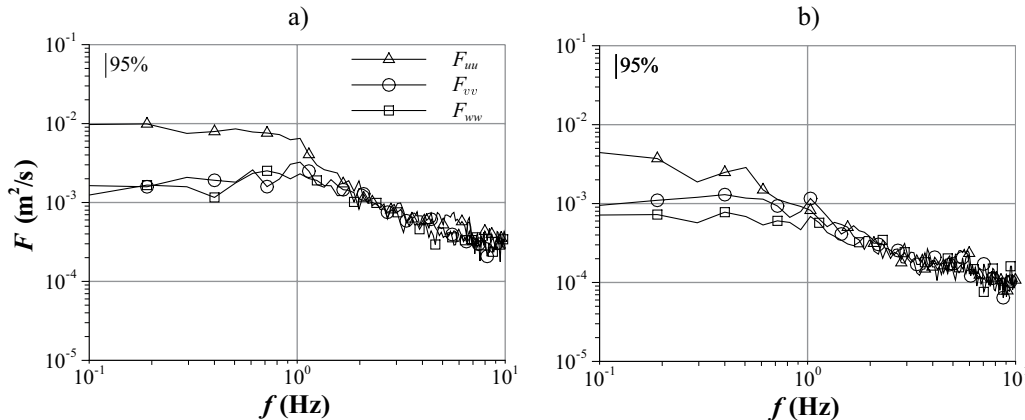


FIGURE 10. Flow velocity spectrum a) behind the plant patch at  $x = 430, z = 85$  mm; b) above the plant patch at  $x = 430, z = 200$  mm.

688 wake spectrum, but there is no indication of a highly periodic component that would  
 689 suggest a Kelvin Helmholtz type instability of the shear layer.

690

#### 4.6. Plant velocity fluctuations and plant-flow coupling

691 The fluctuating movements of aquatic plants are important for several reasons. First,  
 692 plant motion can enhance photosynthetic rate and nutrient uptake through increased  
 693 delivery of light and nutrients to leaf surfaces (Koehl & Alberte 1988; Nikora 2010). Sec-  
 694 ond, plant drag forces (which determine plant survival during high flow periods) may be  
 695 regulated, to some extent, by plant motion. There is some evidence that waving plants  
 696 can experience less drag by aligning themselves with instantaneous velocity streamlines  
 697 ('dynamic reconfiguration', Siniscalchi & Nikora 2013), although in general the reverse  
 698 may also be true. Finally, plant movement can enhance turbulent kinetic energy in the  
 699 plant wake with implications for sediment transport and mixing processes. In the follow-  
 700 ing, we study plant velocity fluctuations extracted from PIV images using the method  
 701 described in section 2.3, to examine the nature of the plant motion and potential inter-  
 702 action mechanisms with the turbulent flow.

703 Measured vertical and transverse plant velocity variance (figure 11a) is found to in-  
 704 crease rapidly approaching the free end of the plant, consistent with the similar measure-  
 705 ments of Siniscalchi & Nikora (2013) for a variety of aquatic plant species in a laboratory  
 706 flume. The shape of the variance distribution reflects the structural dynamics of the  
 707 plant and the turbulent forcing due to the fluctuating viscous and pressure stresses at  
 708 plant surfaces. For simple structures undergoing free vibration (without external forcing)  
 709 analytical solutions to the equations of motion may be obtained to predict the relative  
 710 amplitude (and variance) of vibrations along the structure. The complex geometry of  
 711 aquatic plants and their as yet uncertain biomechanical properties still preclude such  
 712 analysis for the present case without dramatic simplifications. The ratio of transverse to  
 713 vertical plant velocity variance is in the range 1.15 to 1.35 over the resolved plant length,  
 714 quite similar to the corresponding ratio of fluid velocity variance in the wake of the plant  
 715 (1 to 1.2). Further information about the nature of plant velocity fluctuations can be  
 716 obtained from the 2-point correlation function (4.6). Figure 11b indicates that the time  
 717 ( $\Delta t = f_s^{-1} \Delta t_n$ ) corresponding to the maximum in the correlation function is increasing  
 718 with increasing point separation ( $\Delta x = x_d - x_u$ ). This suggests that the characteristic  
 719 plant motion is that of travelling waves rather than standing waves (vibration). These

720 two phenomena are, however, closely related as standing waves can be considered to  
 721 arise from the interference (constructive and destructive) of forward and backward prop-  
 722 agating waves (Graff 1991). Paidoussis (2004), considering slender cylindrical structures  
 723 aligned axially with the flow, indicates that wave propagation rather than vibration is  
 724 typical for long structures. The propagation velocity estimated from the time lag that  
 725 maximises the correlation function (figure 11b) is 0.46 m/s for both  $v_p'$  and  $w_p'$ , which  
 726 is similar to the eddy convection velocity measured in the wake of the plant patch in  
 727 the shear zone (figure 9). The similarity between these two convection velocities suggests  
 728 that the waves propagating through the plant are dominated by the passage of turbulent  
 729 fluctuations (vortices). The plant velocity spectrum (figure 11c,  $x = 309$  mm) indicates  
 730 maximum energy for frequencies around 1 Hz for both vertical and transverse compo-  
 731 nents. The shape of the spectrum resembles that obtained in laboratory experiments  
 732 using the same species of plant (*Ranunculus penicillatus*, Siniscalchi & Nikora 2013) and  
 733 features a significant decay of energy towards both lower and higher frequencies. In com-  
 734 parison, the transverse and vertical components of the fluid velocity spectrum measured  
 735 outside the flow region influenced by the plant (figure 10b) are constant (saturated) for  
 736 frequencies less than 1 Hz. If the plant velocity can be considered as a (linearly) filtered  
 737 response to the fluid velocity, figure 11c in comparison to figure 10b suggests that the  
 738 plant responds optimally to frequencies around 1 Hz (or wavelengths  $\bar{u}/f \sim 0.5$  m, i.e.  
 739 of a similar scale to the patch length or flow depth). This observation may be related  
 740 to Naudascher & Rockwell's (1994) finding that for cylinders aligned axially with the  
 741 flow, each vibration mode of the structure is most efficiently excited by vortices of a cer-  
 742 tain wavelength. Vibration modes for an aquatic plant are, however, yet to be identified.  
 743 Possible mechanisms of flow-plant interaction are further discussed in section 5.

744 Potential correlations between fluid ( $u_j'$ ) and plant ( $u_{ip}'$ ) velocity fluctuations can be  
 745 further examined using the normalised covariance function:

$$R_{0i_pj}(x, z) = \frac{\overline{u_{ip}'(x) u_j'(x, z)}}{\left(\overline{u_{ip}'^2} \overline{u_j'^2}\right)^{0.5}} \quad (4.11)$$

746 with  $i=2, 3$  ( $v_p', w_p'$ ) and  $j=1, 2, 3$  ( $u', v', w'$ ). For  $x = 309$  mm and  $z$  values approaching  
 747 the free surface, the correlation between plant and fluid motion for all components is  
 748 small ( $R_{0i_pj} \sim 0.05$ , figure 11d). The  $R_{0v_pu}$ ,  $R_{0v_pw}$ ,  $R_{0w_pv}$  terms remain small for all  $z$ ,  
 749 but the  $R_{0v_pv}$ ,  $R_{0w_pw}$ ,  $R_{0w_pu}$  terms increase rapidly approaching the top of the plant.  
 750 While it is not surprising to find a correlation between matching velocity components  
 751 ( $R_{0v_pv}$ ,  $R_{0w_pw}$ ) and the cross-component term  $R_{0w_pu}$  through the secondary correlation  
 752  $\overline{u'w'} < 0$ , the narrowness of the correlated range  $\Delta z \sim 30$  mm is unexpected. Given  
 753 the 1 Hz characteristic frequency (figure 11c) of plant velocity fluctuations, we might  
 754 reasonably be looking for characteristic eddy sizes of the order  $\bar{u}/f \sim 0.5$  m and a  
 755 correspondingly larger correlation length. In the following section we consider potential  
 756 flow-plant interaction mechanisms that may help interpret the measured spectra and  
 757 correlation functions.

## 758 5. Flow-plant interactions: concluding remarks

759 Naudascher & Rockwell (1994) identified three general classes of flow-induced vibra-  
 760 tion mechanisms: extraneously induced excitation (EIE), instability induced excitation  
 761 (IIE), and movement induced excitation (MIE). These classifications were developed to  
 762 help identify and analyse the source of vibrations in engineering structures, but they are  
 763 also relevant to the present case of flow-aquatic plant interactions, even if the charac-



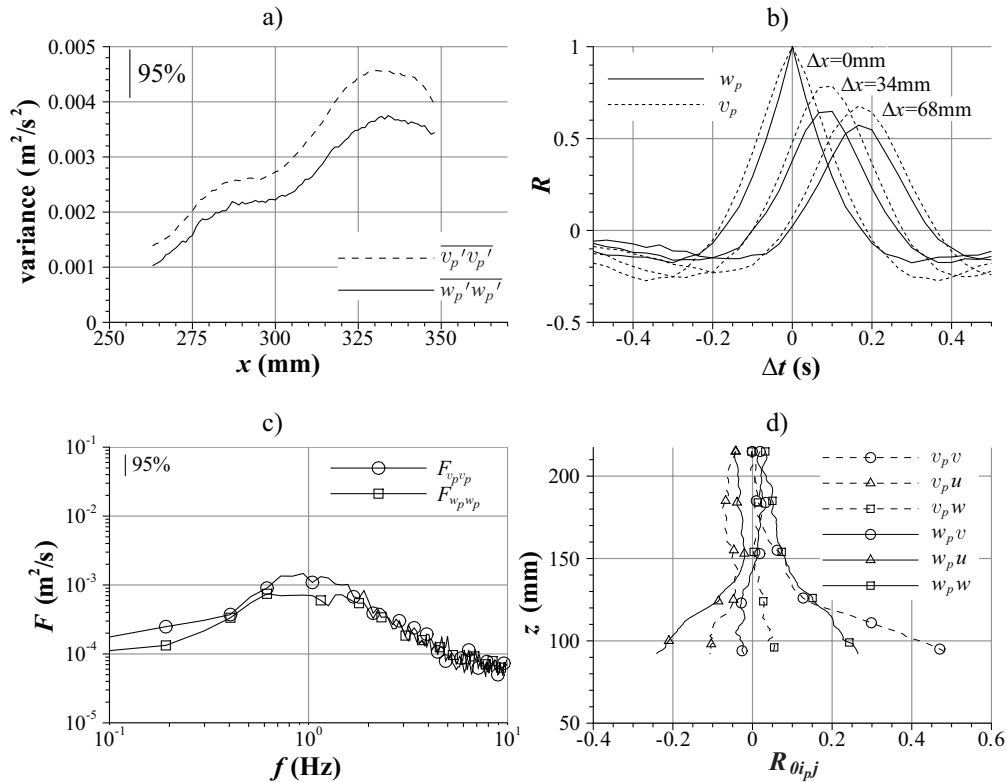


FIGURE 11. Variation of plant velocity variance along the length of the plant (a). Two-point correlation of plant velocity fluctuations (b). Spectrum of plant velocity fluctuations ( $x = 309$  mm, c). Normalised covariance between plant and flow velocity fluctuations ( $x = 309$  mm, d).

764 teristic plant motion is that of a propagating wave (figure 11b) rather than a vibration.  
 765 Extraneously induced excitation relates to structure (plant) motion caused by turbulence  
 766 in the flow, but independent of any local flow instability associated with the presence of  
 767 the plant. Instability induced excitation relates to motion induced by a flow instability  
 768 that appears due to the presence of a structure. For an aquatic plant, this instability  
 769 could, for example, be unsteady flow separation from the plant or a shear layer insta-  
 770 bility. Flow separation at a plant scale seems unlikely based on figure 6, however, there  
 771 is significant turbulent kinetic energy associated with the shear layer in the wake of  
 772 the plant patch (figures 8, 9), and the short range of elevations over which plant and  
 773 fluid velocity fluctuations are correlated (figure 11d) support a contribution of IIE to the  
 774 plant motion. The absence of strong periodicity in the velocity spectrum measured in  
 775 the plant patch wake (figure 10a), however, does not support an instability of the Kelvin  
 776 Helmholtz type. We did observe a weak clustering of energy around 1 Hz (matching well  
 777 the dominant frequency of plant motion, figure 11c), however, this may simply reflect  
 778 the flapping elevation of the shear layer as the plant moves up and down. The relative  
 779 importance of EIE and IIE cannot be confirmed from the present experiments, but this  
 780 could be further investigated in a laboratory environment by, for example, towing plants  
 781 through stationary water to eliminate sources of EIE. Movement induced excitation is  
 782 a self-excited body vibration where the acceleration of a body in a fluid alters the flow  
 783 field in a way that can feed back to the body (via pressure and viscous stresses) to am-  
 784 plify the initial movement. The ‘flutter’ of flags or aircraft wings are examples of MIE.

785 The correspondence between measured convection velocities of plant velocity fluctuations  
 786 and fluid velocity fluctuations in the plant patch wake suggests that the plant velocity  
 787 fluctuations are dominated by the passage of turbulent eddies (either EIE and IIE), and  
 788 MIE seems unlikely in the present case. Flexible cylinders aligned axially with the flow  
 789 (resembling aquatic plant stems to some extent) can exhibit MIE at certain critical flow  
 790 velocities, the dynamics for which have been studied extensively (e.g. Païdoussis 2004;  
 791 de Langre *et al.* 2007). Extension of this type of analysis to an aquatic plant is not yet  
 792 realistic due to the complex and changing plant geometry (reconfiguration) and the lack  
 793 of fundamental knowledge of the coupling between fluid flow and resulting lift and drag  
 794 forces acting on the plant.

795 Further experimentation in the laboratory and in the field is needed to clarify the  
 796 nature of flow-plant interaction mechanisms. In this regard, we have demonstrated that  
 797 the stereoscopic PIV method can be applied in field conditions and should prove to be  
 798 valuable in further study of flow-aquatic plant interactions.

799  
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#### REFERENCES

- 804 ALBAYRAK, I., NIKORA, V., MILER, O. & O’HARE, M. 2012 Flow-plant interactions at a leaf  
 805 scale: Effects of leaf shape, serration, roughness and flexural rigidity. *Aquat. Sci.* **74** (2),  
 806 267–286.
- 807 ASTARITA, T. 2007 Analysis of weighting windows for image deformation methods in PIV. *Exp.*  
 808 *Fluids* **43** (6), 859–872.
- 809 BALACHANDAR, R., BLAKELY, D., TACHIE, M. & PUTZ, G. 2001 A study on turbulent boundary  
 810 layers on a smooth flat plate in an open channel. *J. Fluids Eng.* **123** (2), 394–400.
- 811 BORNETTE, G. & PUJALON, S. 2011 Response of aquatic plants to abiotic factors: A review.  
 812 *Aquat. Sci.* **73** (1), 1–14.
- 813 CALLUAUD, D. & DAVID, L. 2004 Stereoscopic particle image velocimetry measurements of the  
 814 flow around a surface-mounted block. *Exp. Fluids* **36** (1), 53–61.
- 815 CAMERON, S.M. 2011 PIV algorithms for open-channel turbulence research: Accuracy, resolution  
 816 and limitations. *J. Hydro-environ. Res.* **5** (4), 247–262.
- 817 CANTWELL, B. & COLES, D. 1983 Experimental study of entrainment and transport in the  
 818 turbulent near wake of a circular cylinder. *J. Fluid Mech.* **136**, 321–374.
- 819 DAVIDSON, P.A. 2004 *Turbulence: An Introduction for Scientists and Engineers*. Oxford Uni-  
 820 versity Press.
- 821 FOLKARD, A.M. 2005 Hydrodynamics of model *Posidonia oceanica* patches in shallow water.  
 822 *Limnol Oceanogr* **50** (5), 1592–1600.
- 823 FOLKARD, A.M. 2011a Flow regimes in gaps within stands of flexible vegetation: Laboratory  
 824 flume simulations. *Environ. Fluid Mech.* **11** (3), 289–306.
- 825 FOLKARD, A.M. 2011b Vegetated flows in their environmental context: A review. *Proceedings*  
 826 *of the Institution of Civil Engineers: Engineering and Computational Mechanics* **164** (1),  
 827 3–24.
- 828 GARCIA, C.M., JACKSON, P.R. & GARCIA, M.H. 2006 Confidence intervals in the determination  
 829 of turbulence parameters. *Exp. Fluids* **40** (4), 514–522.
- 830 GHISALBERTI, M. & NEPF, H.M. 2002 Mixing layers and coherent structures in vegetated  
 831 aquatic flows. *J. Geophys. Res. C: Oceans* **107** (2), 3–1.
- 832 GRAFF, K.F. 1991 *Wave Motion in Elastic Solids*. Dover.
- 833 GREEN, J.C. 2005 Velocity and turbulence distribution around lotic macrophytes. *Aquatic Ecol-*  
 834 *ogy* **39** (1-2), 1–10.

- 835 HURTHER, D. & LEMMIN, U. 2001 A correction method for turbulence measurements with a  
836 3D acoustic doppler velocity profiler. *J. Atmos. Oceanic Technol.* **18** (3), 446–458.
- 837 HUSSEIN, H.J. & MARTINUZZI, R.J. 1996 Energy balance for turbulent flow around a surface  
838 mounted cube placed in a channel. *Phys. Fluids* **8** (3), 764–780.
- 839 KATIJA, K. & DABIRI, J.O. 2008 In situ field measurements of aquatic animal-fluid interactions  
840 using a self-contained underwater velocimetry apparatus (SCUVA). *Limnol. Oceanogr.:  
841 Methods* **6** (APR), 162–171.
- 842 KOEHL, M.A.R. & ALBERTE, R.S. 1988 Flow, flapping, and photosynthesis of nereocystis  
843 leutkeana: a functional comparison of undulate and flat blade morphologies. *Mar. Biol.*  
844 **99** (3), 435–444.
- 845 DE LANGRE, E. 2008 Effects of wind on plants. *Annu. Rev. Fluid Mech.* **40**, 141–168.
- 846 DE LANGRE, E., PADOUSSIS, M.P., DOAR, O. & MODARRES-SADEGHI, Y. 2007 Flutter of long  
847 flexible cylinders in axial flow. *J. Fluid Mech.* **571**, 371–389.
- 848 LIAO, Q., BOOTSMA, H.A., XIAO, J., KLUMP, J.V., HUME, A., LONG, M.H. & BERG, P.  
849 2009 Development of an in situ underwater particle image velocimetry (UWPIV) system.  
850 *Limnol. Oceanogr.: Methods* **7** (FEB.), 169–184.
- 851 LOMB, N.R. 1976 Least-squares frequency analysis of unequally spaced data. *Astrophys. Space  
852 Sci.* **39** (2), 447–462.
- 853 MAAS, H.G. 1996 Contributions of digital photogrammetry to 3D PTV. In *Three-dimensional  
854 velocity and vorticity measuring and image analysis techniques* (ed. T. Dracos), pp. 191–  
855 208. Kluwer.
- 856 MANSOUR, N.N., KIM, J. & MOIN, P. 1988 Reynolds-stress and dissipation-rate budgets in a  
857 turbulent channel flow. *J. Fluid Mech.* **194**, 15–44.
- 858 NADEN, P., RAMESHWARAN, P., MOUNTFORD, O. & ROBERTSON, C. 2006 The influence of  
859 macrophyte growth, typical of eutrophic conditions, on river flow velocities and turbulence  
860 production. *Hydrol. Processes* **20** (18), 3915–3938.
- 861 NAUDASCHER, E. & ROCKWELL, D. 1994 *Flow-Induced Vibrations: An Engineering Guide*.  
862 Dover.
- 863 NEPF, H.M. 2012 Hydrodynamics of vegetated channels. *J. Hydraul. Res.* **50** (3), 262–279.
- 864 NEZU, I. & NAKAGAWA, H. 1993 *Turbulence in Open Channel Flows*. Balkema.
- 865 NEZU, I. & SANJOU, M. 2008 Turbulence structure and coherent motion in vegetated canopy  
866 open-channel flows. *J. Hydro-environ. Res.* **2** (2), 62–90.
- 867 NIKORA, N. & NIKORA, V. 2007 A viscous drag concept for flow resistance in vegetated channels.  
868 In *Harmonizing the demands of art and nature in hydraulics: Proc. of 32nd Congress of  
869 IAHR*. Venice, Italy, (CD-ROM).
- 870 NIKORA, V. 2010 Hydrodynamics of aquatic ecosystems: An interface between ecology, biome-  
871chanics and environmental fluid mechanics. *River Res. Appl.* **26** (4), 367–384.
- 872 NIKORA, V., CAMERON, S., ALBAYRAK, I., MILER, O., NIKORA, N., SINISCALCHI, F., STEW-  
873ART, M. & O’HARE, M. 2012 Flow-biota interactions in aquatic systems: Scales, mecha-  
874nisms, and challenges. In *Environmental Fluid Mechanics: Memorial volume in honour of  
875 Prof. Gerhard H. Jirka* (ed. W. Rodi & M. Uhlmann), pp. 217–235. CRC Press.
- 876 NIKORA, V. & GORING, D.G. 2000a Eddy convection velocity and Taylor’s hypothesis of ‘frozen’  
877 turbulence in a rough-bed open-channel flow. *Journal of Hydrosience and Hydraulic En-  
878 gineering* **18** (2), 75–91.
- 879 NIKORA, V. & GORING, D. 2000b Flow turbulence over fixed and weakly mobile gravel beds.  
880 *J. Hydraul. Eng.* **126** (9), 679–690.
- 881 NIMMO SMITH, W.A.M., ATSAVAPRANEE, P., KATZ, J. & OSBORN, T.R. 2002 PIV measure-  
882 ments in the bottom boundary layer of the coastal ocean. *Exp. Fluids* **33** (6), 962–971.
- 883 NOBACH, H. & BODENSCHATZ, E. 2009 Limitations of accuracy in PIV due to individual vari-  
884 ations of particle image intensities. *Exp. Fluids* **47** (1), 27–38.
- 885 O’HARE, M.T., HUTCHINSON, K.A. & CLARKE, R.T. 2007 The drag and reconfiguration ex-  
886 perimented by five macrophytes from a lowland river. *Aquat. Bot.* **86** (3), 253–259.
- 887 PAÏDOUSSIS, M.P. 2004 *Fluid-Structure Interactions: Slender Structures and Axial Flow, Volume  
888 2*. Elsevier.
- 889 POPE, S.B. 2000 *Turbulent Flows*. Cambridge University Press.
- 890 PRASAD, A.K. 2000 Stereoscopic particle image velocimetry. *Exp. Fluids* **29** (2), 103–116.

- 891 RAFFEL, M., WILLERT, C.E., WERELEY, S.T. & KOMPENHANS, J. 2007 *Particle Image Ve-*  
892 *locimetry: A Practical Guide*. Springer-Verlag.
- 893 RAUPACH, M.R., FINNIGAN, J.J. & BRUNET, Y. 1996 Coherent eddies and turbulence in veg-  
894 etation canopies: the mixing-layer analogy. *Boundary Layer Meteorol.* **78** (3-4), 351–382.
- 895 SAND-JENSEN, K. 2003 Drag and reconfiguration of freshwater macrophytes. *Freshwater Biology*  
896 **48** (2), 271–283.
- 897 SAND-JENSEN, K. & MEBUS, J.R. 1996 Fine-scale patterns of water velocity within macrophyte  
898 patches in streams. *Oikos* **76** (1), 169–180.
- 899 SCARGLE, J.D. 1982 Studies in astronomical time series analysis. II. Statistical aspects of spec-  
900 tral analysis of unevenly spaced data. *The Astrophysical Journal* **263** (2), 835–853.
- 901 SHAW, R.H., BRUNET, Y., FINNIGAN, J.J. & RAUPACH, M.R. 1995 A wind tunnel study of  
902 air flow in waving wheat: Two-point velocity statistics. *Boundary Layer Meteorol.* **76** (4),  
903 349–376.
- 904 SINISCALCHI, F. & NIKORA, V.I. 2012 Flow-plant interactions in open-channel flows: A com-  
905 parative analysis of five freshwater plant species. *Water Resour. Res.* **48** (5).
- 906 SINISCALCHI, F. & NIKORA, V.I. 2013 Dynamic reconfiguration of aquatic plants and its inter-  
907 relations with upstream turbulence and drag forces. *J. Hydraul. Res.* **51** (1).
- 908 SINISCALCHI, F., NIKORA, V.I. & ABERLE, J. 2012 Plant patch hydrodynamics in streams:  
909 Mean flow, turbulence, and drag forces. *Water Resour. Res.* **48** (1).
- 910 SUKHODOLOVA, T.A. & SUKHODOLOV, A.N. 2012 Vegetated mixing layer around a finite-size  
911 patch of submerged plants: 1. theory and field experiments. *Water Resour. Res.* **48** (10).
- 912 TRITICO, H.M., COTEL, A.J. & CLARKE, J.N. 2007 Development, testing and demonstration of  
913 a portable submersible miniature particle imaging velocimetry device. *Meas. Sci. Technol.*  
914 **18** (8), 2555–2562.
- 915 UBEROI, M.S. & FREYMUTH, P. 1970 Turbulent energy balance and spectra of the axisymmetric  
916 wake. *Phys. Fluids* **13** (9), 2205–2210.
- 917 VOGEL, S. 1994 *Life in Moving Fluids: The Physical Biology of Flow*. Princeton University  
918 Press.
- 919 WERNET, M.P. 2005 Symmetric phase only filtering: A new paradigm for DPIV data processing.  
920 *Meas. Sci. Technol.* **16** (3), 601–618.
- 921 WIENEKE, B. 2005 Stereo-PIV using self-calibration on particle images. *Exp. Fluids* **39** (2),  
922 267–280.
- 923 WILLERT, C. 1997 Stereoscopic digital particle image velocimetry for application in wind tunnel  
924 flows. *Meas. Sci. Technol.* **8** (12), 1465–1479.
- 925 WYGNANSKI, I. & FIEDLER, H.E. 1970 Two-dimensional mixing region. *J. Fluid Mech.* **41** (pt  
926 2), 327–361.
- 927 ZHU, W., VAN HOUT, R., LUZNIK, L., KANG, H.S., KATZ, J. & MENEVEAU, C. 2006 A  
928 comparison of PIV measurements of canopy turbulence performed in the field and in a  
929 wind tunnel model. *Exp. Fluids* **41** (2), 309–318.